

Brachyoxylon cristianicum sp. nov., *Protophylocladoxylon holbavicum* sp. nov., and *Palaeoginkgoxylon* sp. The present paper is based exclusively on the study of the silicified woods in the “Grădinaru Collection”. As many others silicified woods were collected in the last decades in the Holbav and Cristian areas by the authors of the present paper, their xylotomical study will be done in future papers.

GEOLOGICAL SETTING AND STRATIGRAPHIC DATA

The Lower Jurassic Gresten-type deposits with coals are known for a long time in the Holbav and Cristian areas, both located in the Braşov region. According to Săndulescu (1984), the Holbav area is located in the Holbav Unit while the Cristian area is located in the Braşov Unit, both units belonging to the Getic domain that is prolonged in the easternmost part of the Southern Carpathians (Fig. 1).

Holbav area

In the Holbav area, the industrial mining extraction of coal started early in the second half of the 19th century and continued until the late 1960s when it stopped due to limited coal reserves and difficult geological conditions (Bielz, 1858; Meschendorfer, 1860; Hauer & Stache, 1863; Herbich, 1878; Römer, 1878, 1879, 1885; Wachner, 1914, Jekelius, 1923a-b, 1927, 1938; Vîlceanu, 1960; Semaka, 1962d, 1967a; Mateescu, 1964; Năstăseanu *et al.*, 1970; Săndulescu *et al.*, 1984).

Obviously, the Lower Jurassic deposits with coals are poorly appearing as natural outcrops in the Holbav area. The description of the stratigraphic successions and of the lithologies of the Lower Jurassic deposits, such as it has been given by the different authors, are based mainly on data obtained from mining works and from boreholes. The tectonic structure of the Holbav Unit (or the so-called "Holbav Synclinal" in the older literature), interpreted as occurring in a tectonic window by Săndulescu (1967; 1984, figures 83, 86), is very complicated due to the

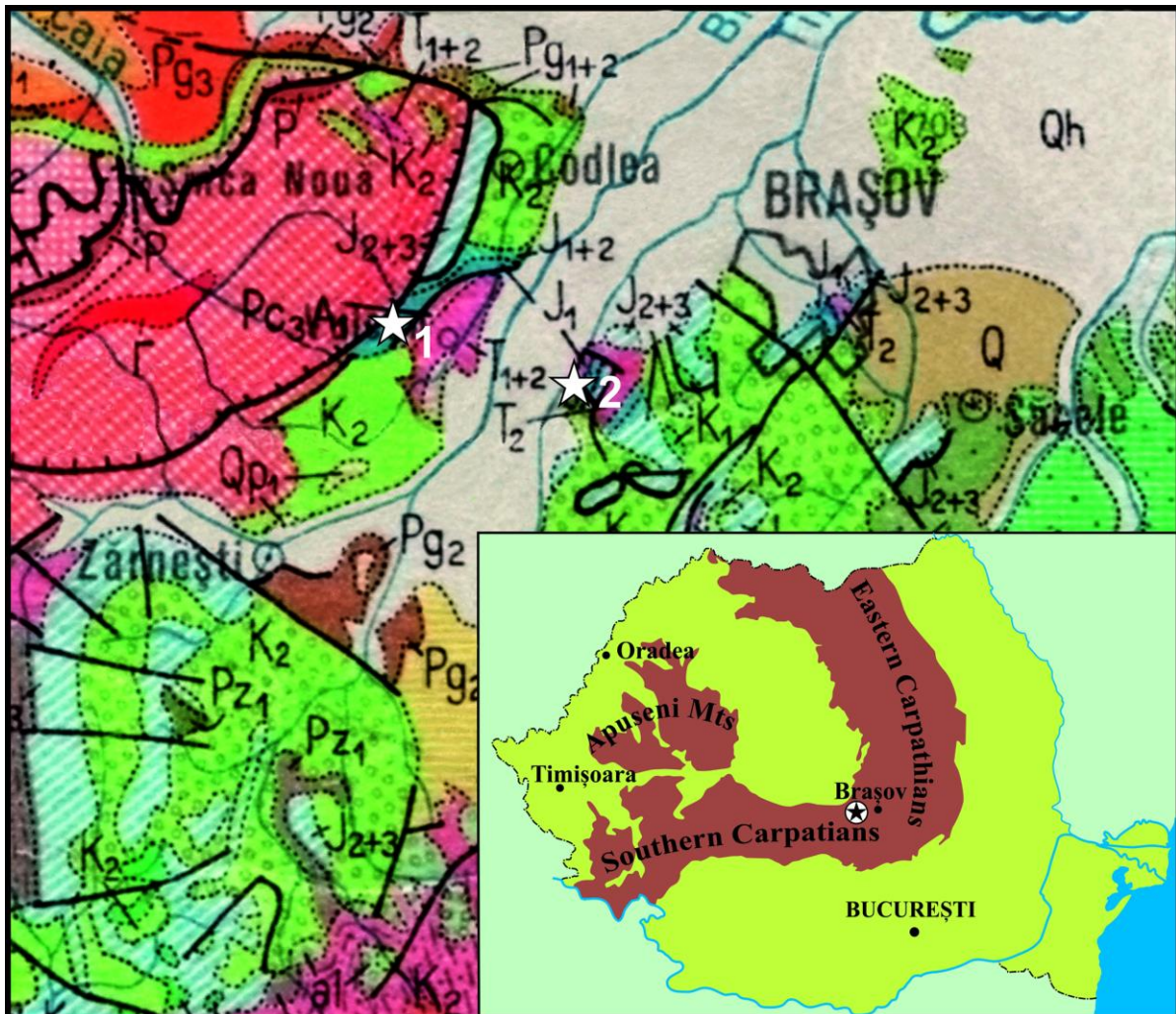


Fig. 1 Location of the Holbav (1) and Cristian (2) areas on the Geological Map of Romania, Scale 1:100000000, ed.1978, Geological Institute of Romania (excerpt). Inset map shows the location of the Braşov region (encircled star), in the easternmost part of the Southern Carpathians.

overthrusting of the Supragetic Unit along the Holbav Fault, and also due to the overthrusting of the Getic Măgura-Codlei Scale and the Braşov Unit.

As a consequence, the stratigraphical and lithological descriptions of the Lower Jurassic deposits in the Holbav Unit given along time by the different authors (Jekelius, 1923a-b, 1927, 1938; Vilceanu, 1960; Manilici & Vilceanu, 1962; Semaka, 1962d, 1967a; Mateescu, 1964; Năstăseanu *et al.*, 1970; Săndulescu *et al.*, 1984; Ştefănuţ & Cibotaru, 1986) are, at most, informative (Grădinaru, in Philippe *et al.*, 2006, and Grădinaru, in Haas *et al.*, 2011). Vilceanu (1960), based mainly on data from mining works and boreholes, identified in the Lower Jurassic deposits with coals from the Holbav area (Fig. 2) four lithological complexes, as follows:

- the lower productive complex,
- the effusive-pyroclastic complex (bearing silicified woods),
- the upper productive complex, and
- the sandstone complex.

Patruluiş *et al.* (1980) and Săndulescu *et al.* (1984) provided a general lithological description of the Lower Jurassic coaly deposits from the Holbav area, with three sequences:

- the lower sequence, mainly with clays and sandstones bearing coaly clays and coals, occasionally interlayered by reddish-cherry clays;
- the median sequence, volcano-sedimentary, with rhyolitic and trachytic pyroclastites, tuffs, tuffaceous rocks and eruptive rocks such as ignimbrites, trachytes, and basaltic rocks, subordinately having intercalations of clays, bituminous claystones, coaly clays and thin-bedded sandstones.
- the upper sequence, mainly with clays and sandstones, bearing coal intercalations and refractory clays.

All of the above lithostratigraphic schemes are based mainly on data from mining works and boreholes, and not on objective data coming from mapping and definition of stratigraphic relationships in outcrops. The strong tectonic deformation, the poor outcrops and the complex depositional architecture that is characteristic for the alluvial and the coal-bearing lacustrine deposits (Reading, 1996), all of these make difficult to establish the stratigraphic succession of the Lower Jurassic deposits in the Holbav Unit. On the other side, the unevenly distribution of the effusive rocks and pyroclastites in the Lower Jurassic coaly deposits in the Holbav area (Săndulescu *et al.*, 1986) makes difficult the distinguishing among the lower and the upper sequences with coals.

Dragastan & Popa (1997) informally grouped the Lower Jurassic deposits with coals from the Holbav area under the name of the "Holbav Formation", with two members, the "Concordia Member" (early Hettangian), and the "Pârâul Crucii Member" (middle Hettangian - Sinemurian - Pliensbachian), respectively.

On his turn, Czier (1999a) introduced the "Codlea-Vulcan Formation" (Hettangian-Pliensbachian), with three members, the "Codlea Conglomerate Member" (Hettangian pro parte), the "Vulcan Sandstone Member" (Hettangian pro parte-Sinemurian), and the "Holbav Clay Member" (Pliensbachian). It should be underlined that the both lithostratigraphic schemes are informal, because none of them is following the principles of stratigraphic classification

defined by the International Stratigraphic Guide (Murphy & Salvador, 1999). None of the above-named lithostratigraphic units are defined based on type sections or stratotypes, whereas the indicated ages are not based on reliable biostratigraphic data.

The only firm biostratigraphic data for dating the Lower Jurassic deposits with coals in the Holbav area have been provided by Antonescu (1973) and Antonescu (in Săndulescu *et al.*, 1984) based on palynological investigations. In the upper half of the upper sequence defined by Săndulescu *et al.* (1984), occurring on the Bisericii brook, a Toarcian microflora has been identified, with *Matonisporites crassiangulatus* (Balme) Dettmann, *Calialasporites* spp., *Neoraistrikia samuelssoni* Tralau, *Sestrosporites pseudoalveolatus* Dettman, *Nannoceratopsis gracilis* Alberti, *N. plegas* Drugg, *Leptolepidites major* Couper, and other genera and species of Early Jurassic age. From the base of the upper sequence, just above the volcano-sedimentary sequence, a Pliensbachian microflora has been identified, with *Ischyosporites variegatus* Schultz (extremely frequent), cf. *Manumia irregularis* Pocock and cf. *Cingulatisporites*. A coaly-siltic intercalation, located in the lower part of the volcano-sedimentary sequence, delivered a rich Pliensbachian assemblage with acritarchs of *Micrhystridium* - type, which is correlatable with the acritarch assemblage of the Pliensbachian spathic sandstones from the Cristian area. It is assumed that the lower sequence with coals from the Holbav area may be assigned to the Hettangian(?)-Sinemurian interval.

In the mid 1970s, Grădinaru (unpublished data) collected an early Toarcian dactylioceratid ammonite specimen and also montlivaltid corals in the fine-grained micaceous sandstones of the upper sequence, interlayered by coaly deposits, occurring in the quarry for extraction of refractory clays, which fully confirms the palynological dating by Antonescu (1973) and Antonescu (in Săndulescu *et al.*, 1984), and argues also for the marine ingression. On the other side, the finding of the middle Toarcian ammonite *Hildoceras sublevisoni* Fucini, in a dark black mudstone (Grădinaru, unpublished data) documents that the Early Jurassic sedimentation continued in the Holbav area, in fully marine facies in the middle Toarcian.

Plant remains, preserved as foliage in the Lower Jurassic coaly deposits in the Holbav area have been mentioned early by Andrae, 1852; Stur, 1860a-b; 1872; Meschendörfer, 1860; Hauer & Stäche, 1863; Herbich, 1878; Römer, 1878; Toula, 1911; Krasser, 1922. Numerous modern studies on the "Liassic" flora from Holbav area have been published by Semaka (1955, 1956, 1957, 1958, 1962d, 1965b, 1967a, 1971, 1972) and Semaka & Givulescu (1965). Also, Givulescu (1992c, 1997b) and Givulescu & Czier (1990) have contributed to the knowledge on the Early Jurassic flora in the Holbav area.

Data on the plant occurrences and on coal petrography in the Holbav area have been given by Mateescu (1964) and Năstăseanu *et al.* (1970). In the recent decades, contributions to the Early Jurassic flora in the Holbav area have been given by Dragastan & Popa (1997), Popa (1997a, 1998, 2000b-c) and Popa & van Konijnenburg-van Cittert (1999, 2006). Dragastan & Popa (1997) and Czier (2016b) advanced phytostratigraphic schemes for the Early Jurassic foliar material identified in the Holbav

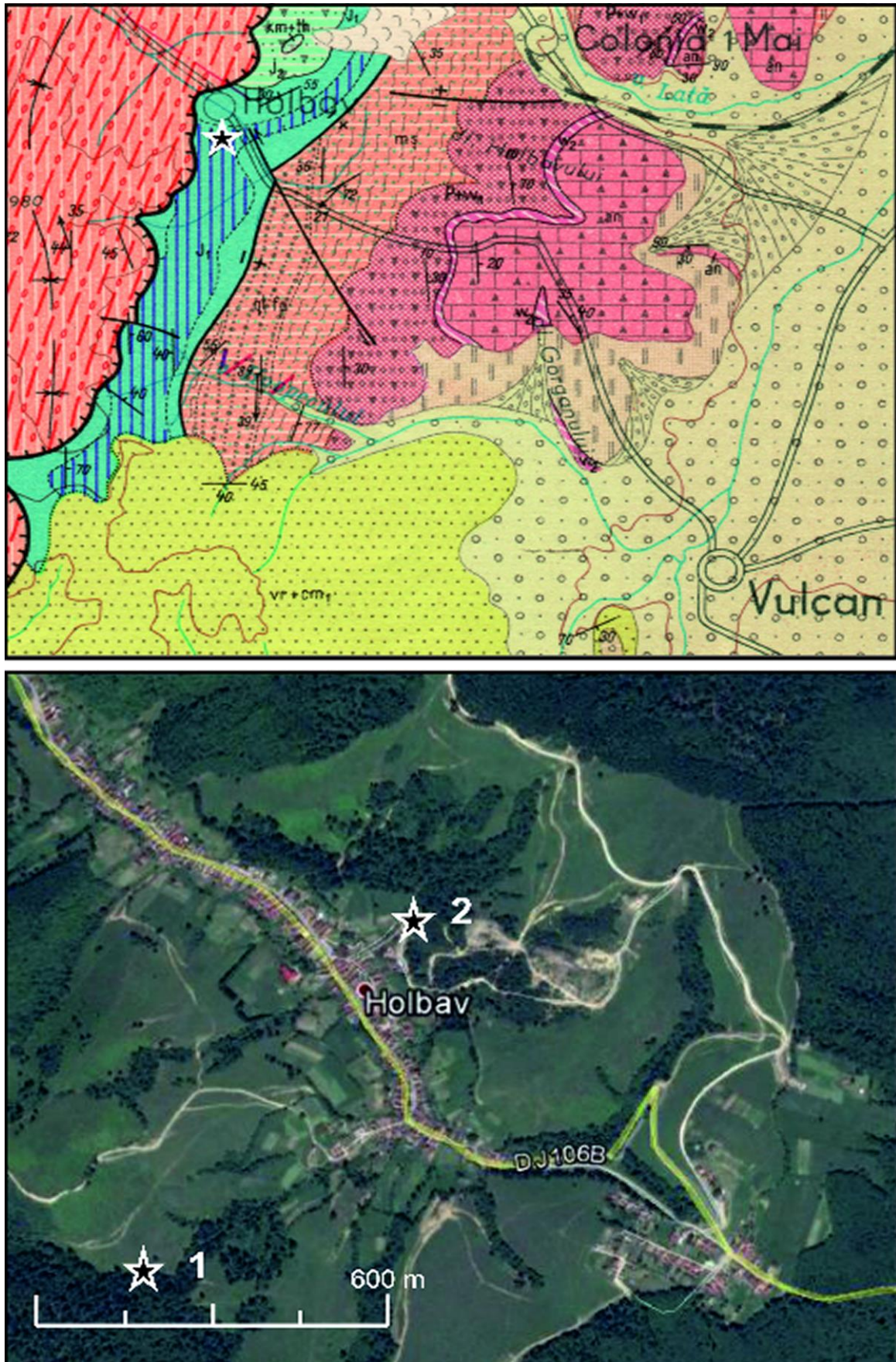


Fig. 2 Location of the Holbav area on the Geological Map of Romania, 1:50000, Sheet 110b Zărnești (excerpt), Geological Institute of Romania (excerpt), and Google Earth image (2018 CNES/Airbus) showing the location on the Maiului brook (star 1) and on the Bisericii brook (star 2) where the Early Jurassic woods were collected. Symbols: P-w1, – Permian-lower Werfenian (Lower Triassic), w2 – upper Werfenian (Lower Triassic); an – Anisian (Middle Triassic).

area. Once more, it should be mentioned that the most part of the Early Jurassic flora described from the Holbav area has been collected from the mining dumps, and thus they do not have a reasonable or reliable stratigraphic control.

With regard to the wood remains in the Lower Jurassic coaly deposits in the Holbav area, it is stressed that Vilceanu (1960) and Manilici & Vilceanu (1962) for the first time mentioned the presence of silicified wood trunks, reaching till 0.70 m in thickness, in the effusive-pyroclastic complex.

In the early 1970s, Grădinaru collected numerous wood trunks, of various sizes, in the middle section of the Maiului brook, and also a few specimens in the middle section of the Bisericii brook (Fig. 2). As concerns the stratigraphic position of the wood trunks, at least in the Maiului brook, there are evidences for their occurrence in the lahar deposits of the uppermost part of the effusive-pyroclastic complex.

Cristian area

In the Cristian area, the mining extraction of coal started in the late 1870s and stopped in the early 1920s (Herbich, 1878; Wachner, 1914; Jekelius, 1923a-b), while the mining extraction of refractory clays continued until the early 1980s. Data on the occurrence and stratigraphy of the Lower Jurassic coaly deposits in the Cristian area have been published by Herbich (1878), Jekelius (1914, 1915a, 1923a-b, 1938), Wachner (1914), Patruleus (1961), Săndulescu (1964), Kusko *et al.* (1979), Patruleus *et al.* (1980), Săndulescu *et al.* (1984). The stratigraphic succession of the Lower Jurassic deposits in the Cristian area is substantially different than that in the Holbav area (Jekelius, 1923a-b, 1938; Patruleus *et al.*, 1980; Săndulescu *et al.*, 1984).

It starts with a coaly sequence with refractory clays, the Hettangian-Sinemurian age of which is fairly demonstrated by the invertebrate fauna (bivalves and ammonites) identified by Jekelius (1914, 1915b) and Popa (in Patruleus *et al.*, 1980). Antonescu (1973), and Antonescu in Patruleus *et al.* (1980) and in Săndulescu *et al.* (1984), identified palynological assemblages that demonstrated the Hettangian-Sinemurian age of the coaly sequence in the Cristian area. The Hettangian age is demonstrated by the occurrence of *Rogalskiasporites cicatricosus* (Rogalska) Schulz, *Convolutispora microrugulata* Schulz, *C. microfoveolata* Schulz, *Tigrisporites rhaeticus* (Schulz) Schurman, *T. rudis* (Leschik) Klaus, *Triancoraesporites ancorae* (Rheinhardt) Schulz, *Auritulinasporites triclavus* Nilsson and *Kyrtomisporites* sp. The Sinemurian age is demonstrated by a very rich assemblage of spores and pollens, including *Dictyophyllidites harrisi* Couper, *Cyathidites australis* Couper, *C. minor* Couper, *Concavisporites juriensis* Balme, *Deltoidospora* sp., *Osmundacidites wellmanii* Couper, *Lycopodiacidites cerebriformis* (Naumova ex Bolchovitina) Semenova, *Punctatosporites scabratus* (Couper), *Cerebropollenites macroverrucosus* (Thiergart) Schulz, *Chasmatosporites apertus* (Rogalska) Nilsson, *Monosulcites minimus* Cookson, *Corollina meyeriana* (Klaus) Venkatachala & Góczán, *Classopollis torosus* (Reissinger) Balme, emend. Morbey, *Bennettitinaepollenites bitorosus* Bóna, *Vitreisporites pallidus* (Re-

issinger) Nilsson, *Obtusisporites junctus* (Kara-Murza) Pocock, *Lycopodium* spp.

On this basis it is demonstrated that the coaly sequence from Cristian is correlable with the lower coaly sequence from the Holbav area. The coaly sequence from Cristian has the most extended distribution on the left side of the Căldării valley and in the regions of the Fabricii and Joaderului valleys. Czier (1999a) introduced the "Cristian Formation" for the coaly sequence from the Cristian area, but it is an informal lithostratigraphic unit because no type section or stratotype is indicated. On the other hand, it is not possible owing to the poor outcropping conditions of the coaly sequence. The coals were extracted only in the Căldării valley, while the refractory clays were extracted in the regions of the Fabricii and Joaderului valleys. Upwards, the succession of the Lower Jurassic deposits from the Cristian area (Fig. 3) is developed in fully marine facies, with Pliensbachian spathic sandstones and Toarcian fine-grained micaceous sandstones (Jekelius, 1914, 1915a, 1938; Săndulescu, 1964; Patruleus *et al.*, 1980).

Thus, when the Lower Jurassic sedimentary successions in the Holbav and Cristian areas are compared, it is to underlie that the coaly deposition in the Holbav area lasted until the lower Toarcian, while in the Cristian area the coaly deposition is limited only to the Hettangian-Sinemurian time interval. It means that while the marine invasion in the Holbav area happened later, in the early Toarcian, in the Cristian area the marine influence happened much earlier and it is demonstrable by the presence of intercalations of siltstones with Sinemurian marine fauna in the coaly sequence (Jekelius, 1915b; Patruleus *et al.*, 1980).

Another marked difference among the Holbav and Cristian areas is represented by the missing or a more reduced and only the local occurrence of the effusive-pyroclastic rocks in the Cristian area. Only greenish ignimbritic rocks were found in the upstream section of the Căldării valley, and in the excavation works for building the PECO oil repository from Cristian (Grădinaru, unpublished data).

Data on the Early Jurassic flora from the Cristian area have been published early by Meschendörfer (1960), Hauer & Stache (1863), Stur (1960a-b, 1872), Herbich (1878), Toula (1911) and Krasser (1922) and these are based on plant remains coming from the mining works in the Căldării valley. Semaka (1965b) and Popa (1998, 2000a-c) discussed on the Early Jurassic flora from the Cristian area based on historical materials, and identified them as being of Sinemurian age.

Based on plant remains collected by Grădinaru in the early 1970s from the dump rocks coming from mines extracting refractory clays in the upstream of the Fabricii valley, Popa (1998) described a new species of *Pachypteris*: *P. grădinarui* Popa.

Czier (2016b) also agrees the Sinemurian age, and assigned the foliage material described from Cristian to his "Anomozamites marginatus Biozone".

Also, in the early 1970s Grădinaru collected some fragments of petrified wood from the right tributary of the Fabricii valley, just upstream of the former charging station of refractory clays, and originating probably from the dumps of the old mine galleries (Fig. 3).

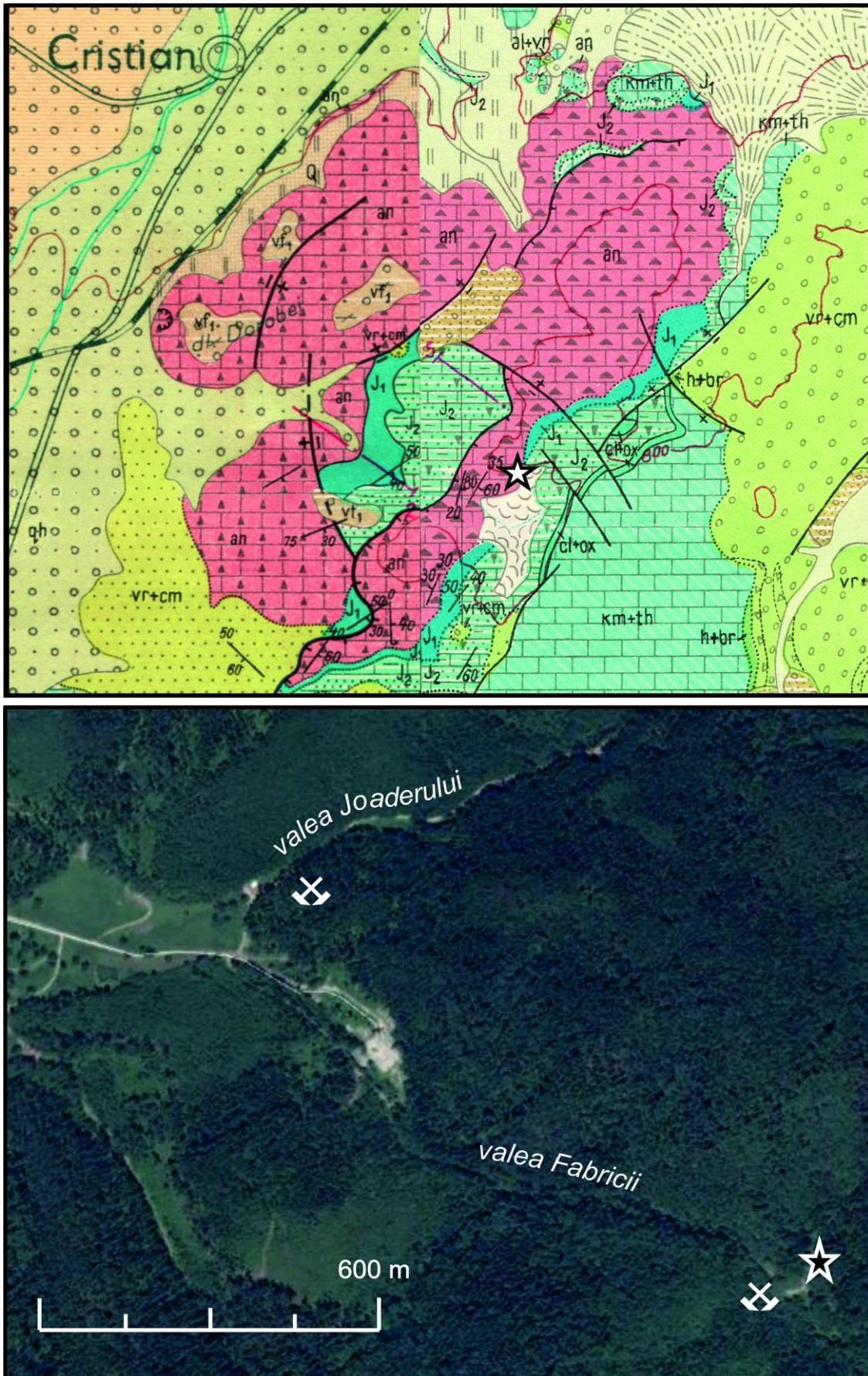


Fig. 3 Location of the Cristian area on the Geological Map of Romania, 1:50000, Sheet 110b Zărnești (excerpt) combined with Sheet 111a Brașov (excerpt), Geological Institute of Romania, and Google Earth image (2018 CNES/Airbus) showing the location on the Valea Fabricii brook (star) where the Early Jurassic woods were collected. Symbols: an – Anisian, Middle Triassic; J1 – Lower Jurassic; J2 – Middle Jurassic; x - abandoned mines.

MATERIAL AND METHODS

A rich fossil wood material was collected in the early 1970s by Grădinaru from the Lower Jurassic coaly deposits in the Holbav and Cristian areas. The most part of the wood material was collected in the mid section of the Maiului brook (Google Earth coordinates: 45°39'12.74" N; 25°22'51.02" E, elev. 776 m), which is a left tributary of the Holbav valley, reaching it in the center of the Holbav village.

Apparently, the material is stratigraphically located in the uppermost part of the effusive-pyroclastic sequence of the Lower Jurassic deposits of the Holbav area, which on palynological data is dated as Pliensbachian in age. A small number of specimens were collected in the Cristian area, on a left tributary reaching the Fabricii valley in its upper section (Google Earth coordinates: 45°36'09.95" N; 25°30'13.23" E, elev. 735 m), all coming from the Hettangian-Sinemurian coaly sequence.

This collection of petrified woods, from which 33 samples were included in the present palaeoxylotomical study, is hosted now by the National Geological Museum, in Bucharest, under the name of "Grădinaru Collection". The most part of the wood material is represented by trunk fragments, sometimes partially charcoaled and included in the lahar deposits where, subsequently, have been silicified. An additional rich material was collected by the authors of this study in the last decades, from the Holbav and Cristian areas, and this will be studied in the future.

The samples of petrified wood were firstly macroscopically studied and the directions of standard oriented sections (transversal, tangential and radial) were carefully marked. In the next stage, following the petrologic method, thin sections were cut and realized (excluding the mounting of lamella, replaced with transparent laquer, even nail polish), in the specialized laboratory of the Geological Institute of Romania, in Bucharest.

The sections have been studied under a microscope with transmitted light (NIKON - Eclipse E400). The micropotos were realized with a video-camera "EverFocus". The plates with microphotos were realized with the help of "PaintShopPro" program combined with "StarOffice5.2"-Presentation.

During the microscopic study a description of the xylotomical details with possible taxonomical value was done for every specimen in all the three standard oriented sections. After all these, in a chapter of "affinities and discussions" we cited other studies and we tried to compare our studied structures with other xylotomically described fossil forms already published, in order to identify the studied fossil wood as a known or new taxon.

The xylotomic description generally followed the nomenclatural code for gymnosperm woods of IAWA Committee (2004) and Ruffinatto *et al.* (2015) and the genera delimitation by Philippe & Bamford (2008) and also, the suprageneric taxonomy of Christenhusz *et al.* (2011) and Chase & Reveal (2009).

PALAEOXYLOTOMICAL STUDY

Gymnospermae

Division **Pinophyta** Cronquist, Takht. & Zimm. ex

Reveal, 1996

Order **Pinales** Gorozhankin, 1904

Genus **Agathoxylon** (Hartig) Philippe, 1995

Agathoxylon holbavicum sp. nov.

Fig. 4 a-i.

The studied material

The studied material is represented by a decimetric-sized piece of silicified wood collected from Holbav locality area (Braşov County), on Maiului brook, in Lower Jurassic continental pyroclastic deposits considered of Pliensbachian age. The studied sample is a trunk or thick-branch fragment with decimetric size, dark to black color and by magnifying glass, or even by naked eye, a regular fibrous structure, suggesting a conifer wood. The studied specimen with the field-number 1041 belong to "Grădinaru Collection", and is deposited in the collections of the Geological Institute of Romania (G.I.R. Collections), hosted by the National Geological Museum, Bucharest, under the inventory number 27611.

Microscopic description

The studied specimen presents *s e c o n d a r y* w o o d with tracheidoxylic structure, with relatively distinct growth-rings boundaries and devoid of any normal or traumatic resin canals.

T r a c h e i d s - with polygonal cross-section with slightly rounded corners, often determining small intercellular spaces. With polygonal rounded lumina unequally sized with radial/tangential diameters of 20-30/20-38 µm and thick walls, of 8-13 µm (double wall), thinner in early-wood: 5-8 µm the double-wall. There are 1-9 radial regular rows of tracheids between two successive rays and the density is (675)1155-1650 tracheids per mm². Tangentially seen, the tracheids are unpitted. The radial pitting is typical araucarian, as uniseriate vertical rows of contiguous hexagonal pits and as biseriate rows with alternate contiguous also hexagonal bordered pits with a diameter of 17-20 µm, and oblique elliptic to vertical apertures of 7.5/2.3 µm in diameters. There are no crassulae or helical thickenings.

A x i a l p a r e n c h y m a - is absent.

M e d u l l a r y r a y s - are thin, rectilinear and are constituted of quadrangular cells with smooth horizontal walls. Tangentially the rays have 1-19 cells in height, i.e. 20-380 µm high. Sometimes the taller rays are locally biseriate, so often they appear as biseriate rays with long endings. The ray cells are polygonal slightly rounded and flattened, unequal in size and relatively thick walled: of 2.5-3.1 µm the simple wall. The ray density is 8-14 rays per horizontal tangential millimeter. Radially the rays are homocellular, cells all procumbent of 20-30 µm high, moderately thick-walled: 5-6.2 µm the double wall. Indentures were not observed. Within the marginal rows the ray-cells are higher, of 30-40(-45) µm, with their outer wall slightly wavy. The typical araucarioid cross-fields have 1-6(-9) oculipores, cupressoid (tending to araucarioid) pits, slightly rounded or hexagonal, of 9-13 µm in diameter, with round or short elliptic tilted apertures to vertical, of 3.5-5.5 µm. In the cross-fields with body ray-cells the pits are 1-3 arranged in one horizontal row, or 2-

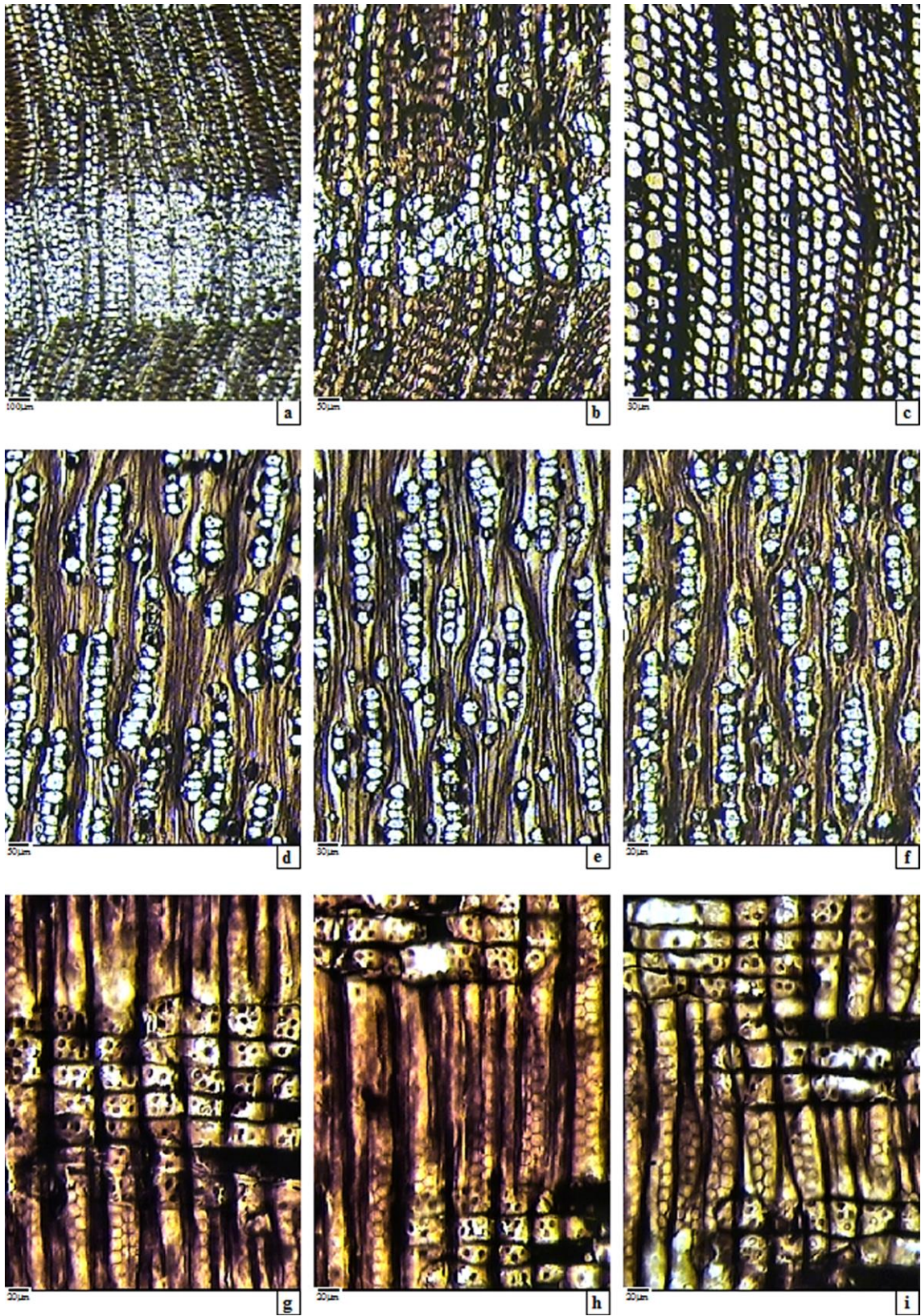


Fig. 4 *Agathoxylon holbavicum* sp. nov. - holotype - inventory no. 27611 (field no. 1041). **a-c** Cross section - tracheids polygonal with rounded corners, and linear rays. **d-f** Tangential section - uniseriate rays with biseriate storeys with polygonal ray-cells, with lateral empty spaces; unpitted tracheids. **g-i** Radial section - araucarian radial pitting on tracheids, alternate hexagonal pits, araucarioid cross-fields with 2-3 superposed rows of hexagonal pits.

6 pits on 2 rows, alternately arranged, or slightly irregular. In the marginal fields they are more numerous, alternately arranged, or slightly irregular, in 2-3 horizontal superposed rows. Often, isolated ray-cells or rows of ray-cells in the body or even marginal cells are full of resin content, visible also in tangential view.

Affinities and discussions

The studied specimen presents a combination of xylotomic characters similar to the extant araucariaceous conifers, consequently a comparison with this type of wood structure was performed and we observed similitude with *Araucaria* and *Agathis* (see Greguss, 1955).

Many fossil woods with tracheidoxyl structure collected from the Palaeozoic and Mesozoic formations have been recorded as species of *Araucarioxylon* and *Dadoxylon*, genera whose status is not legitimate in all cases and thus they have been invalidated (see Philippe, 1993, 1995; Philippe & Barbacka, 1997; Rößler *et al.*, 2014). Thus, for this type of wood was kept the genus *Agathoxylon* Hartig, validly described and published and not yet typified, even if the type material is probably lost (Philippe, 1993).

Other genera as *Dammaroxylon* Schultze-Motel and *Pseudagathoxylon* Greguss are considered as taxonomic synonym of *Agathoxylon*, respectively of *Simplicioxylon* Andreaszky which is a valid genus (see Philippe & Bamford, 2008).

Using the identification key proposed by Philippe (1993) and Philippe & Bamford (2008) for tracheidoxyls with araucarioid cross-fields, we found that the here studied specimen fits to *Agathoxylon* Hartig. The genotype, *A. cordaianum* Hartig, presents similar features having araucarian radial pitting, araucarioid cross-fields and sometimes axial parenchyma. Although the original diagnosis of the genus *Agathoxylon* Hartig (Hartig, 1848; Philippe & Bamford, 2008) includes the presence of axial parenchyma as a diagnostic character (a feature absent in several specimens referred to *Agathoxylon*), Philippe (1995) reformulated the original diagnosis proposed by Hartig (1848), allowing the inclusion of woods with and without axial parenchyma.

This taxonomic decision is adequate since a high number of genera have already been proposed, and erecting another one for including *Agathoxylon*-like woods without axial parenchyma would lead to a more confuse taxonomy.

- We take into account that many European Cretaceous species of *Araucarioxylon*, *Dadoxylon* (*Araucarioxylon*), or *Dadoxylon* (see Schultze-Motel, 1962, 1966), only some of them equivalent to *Agathoxylon*, it is difficult to compare our material with them, but since a serious revision of them is missing. Also, in some recent partial revisions of the Cretaceous woods from Hungary, Italy, France and Austria, it is shown that many of them come from the Lower Cretaceous, or are not tracheidoxyls (Philippe, 1995; Philippe & Barbacka, 1997).

- Some Cretaceous similar tracheidoxyls come also from the Carpathian area (Apuseni mts.): *Agathoxylon* sp. (described initially as *Araucarioxylon* sp.) of Petrescu & Nuțu (1971), *Agathoxylon ultimum* and *A. formosum*, the last described as species of *Dammaroxylon* (by Iamandei

& Iamandei, 2004), all of them presenting the typical araucarian combination of xylotomic characters.

Numerous other Jurassic–Early Cretaceous fossil-woods from Mid-Eastern Europe described as *Agathoxylon* (or equivalents that were reattributed to *Agathoxylon*) are cited by Philippe *et al.* (2006) as follows:

- *Agathoxylon agathiforme* (Kedves) Philippe et Barbacka, (in Philippe & Barbacka, 1997) and *A. parenchymatosum* (Greguss) Philippe, Zijlstra et Barbacka (in Philippe *et al.*, 1999), both of them from the upper Liassic of Hungary.

- *Agathoxylon desnoyersii* (Lemoigne) Philippe, 1995 from the middle Callovian of Poland, initially described as *Araucarioxylon* sp. by Reymanówna (1956), redescribed by Lemoigne as a species of *Brachyphyllum* and now considered only as a junior taxonomical synonym (Philippe *et al.*, 2018).

- *Agathoxylon pannonicum* (Greguss) Barale, Barbacka et Philippe, from the Aptian or Albian of Hungary (see Greguss, 1952; Barale *et al.*, 2002).

Numerous specimens having the typical xylotomic characters: araucarian radial pitting on biseriate rows, contiguous pits, rarely spaced, and only occasionally as uniseriate rows. The cross-field pits are araucarioid, usually as 1-2 cupressoid oculipores arranged in araucarioid manner in the cross-fields were described as *Agathoxylon* sp. either as new data or reattributed to *Agathoxylon*, come from a large area:

- From the Bajocian and Barremian of Bulgaria (Philippe *et al.*, 2006);

- From the Liassic of Hungary (Philippe & Barbacka, 1997; Philippe *et al.*, 1999);

- From the Rhaeto–Liassic and from the Bathonian of Poland (in Philippe *et al.*, 2006);

- From the Bathonian of Romania (Philippe *et al.*, 1999),

- From the late Cretaceous of Spain (García Esteban *et al.*, 2006);

- From Upper Cretaceous (upper Maastrichtian), Dutlu Formation of Turkey (Kutluk *et al.*, 2012);

- Mencl *et al.* (2013) discuss in their paper the occurrence, the anatomical features and the taxonomy of numerous silicified woods from the Late Paleozoic Basins of the Czech Republic attributed to *Agathoxylon*.

- Philippe *et al.* (2015) reattributed an old taxon of Unger described as *Pinites* from the Keuper sandstones in Atteisdorf, Germany as *Agathoxylon keuperianum* nov. comb.

- Also Philippe *et al.* (2018) studying an old collection of Cretaceous fossil woods from France ("Lignier collection"), reattributed some badly identified woods also to *Agathoxylon* as new species or new combinations (*A. crasseradiatum*, *A. tranchantii*, or simply to *Agathoxylon* sp.).

It is obvious that this type of coniferous genus was widespread on the Earth since it was described from already all the continents. Thus, many species were described from the Middle to Upper Cretaceous of Far-East Asia and we quote: *A. tankoense* Stopes & Fujii (initially described as *Dadoxylon tankoense* (Stopes et Fujii) Shimakura (in Philippe, 1993), then an *Agathoxylon kiiense* (Ogura) Oh *et al.*, *A. togeumense* Oh *et al.*, *A. byeon-gpungense* (Kim *et al.*) Oh *et al.*, (in Schultze-Motel, 1966 and in Oh *et al.*, 2011).

- From Africa, we cite a species (*Agathoxylon lifiyii*) described from the Upper Cretaceous of Egypt (Karga Oasis) by Youssef *et al.* (2000) having typical generic features, and also two Permo-Triassic species (*A. africanum*, *A. karoensis*) described by Bamford (2016), in which the araucarian radial pitting is preserved as "Steinkerne" that mimic the mixed type of pitting (see Bamford, 2016), details not observed in our specimen.

- From South America, from the Middle Jurassic of Argentina an *Agathoxylon matildense* with distinct growth-rings, uniseriate pitting on the radial tracheidal walls with contiguous round pits and cross-fields with 4-5 cupressoid pits irregularly arranged and very low rays was described by Zamuner & Falaschi (2005) and other quasi-similar four species by Kloster & Gnaedinger (2017), all of them different of our specimen.

- Also from the Aptian of Patagonia, Vera & Césari (2012) described an *Agathoxylon*, a wood characterized by araucarian tracheid pitting, predominantly uniseriate, rarely biseriate arranged, alternate to opposite and by araucarioid cross-field pits, typical for Araucariaceae. It is true that this type of wood appears also to the Pteridospermales and Cheirolepidiaceae, and the older *Agathoxylon* could represent stems of various cordaitalean - conifer origin (Philippe *et al.*, 2004a). In this respect must be remind that in our specimen description was observed that the taller rays are locally biseriate, often with aspect of biseriate rays with long endings, and also that often ray cells are full of resin content.

- From the Paleocene and Eocene of Western Antarctica some *Agathoxylon* species were described suggesting even possible austral Araucariaceae-dominated forests (Pujana *et al.*, 2014; 2015; Mirabelli *et al.* 2018).

After this discussion of the widespread spread of this Mesozoic genus and taking into account the opinion of García Esteban *et al.* (2006) on the difficulty of distinguishing species of this genus which is species plethoric and the distinction of species is very difficult, so they preferred to describe their material only at generic level, we still believe that we are faced with a new species of *Agathoxylon* that lived within Carpathian area at least during Early Jurassic.

Thus, the studied specimen is characterized by typical araucarian pitting in uniseriate and biseriate-alternate arrangement of contiguous hexagonal pits with oblique-elliptic apertures and typical araucarioid cross-fields with numerous polygonal pits alternately arranged on 1-3 superposed horizontal rows. We named this new species *Agathoxylon holbavicum* sp. nov., remembering the name of the original area (Holbav) from where the studied material was collected. The specimen with inventory no. 27611 (field number 1041) was designated as holotype hosted now in G.I.R. Collections, in the National Geological Museum, Bucharest.

Diagnosis of the new species *Agathoxylon holbavicum* sp. nov.:

Tracheidoxyl with typical araucarian cross-section, with distinct growth rings, without resin canals or axial parenchyma. Tangential tracheidal walls unpitted, radial pitting typical araucarian, uniseriate and biseriate alternate, with contiguous, hexagonal pits with oblique elliptic to vertical apertures. Rays - uniseriate, locally biseriate,

homocellular, marginals higher, with outer wall slightly waved, often with resin content. Cross-fields typical araucarioid with 1-9 oculipores as araucarioid pits alternately arranged on 1-3 horizontal rows, hexagonal or slightly rounded, with short elliptic tilted apertures to vertical.

Genus *Brachyoxylon* Hollick & Jeffrey, 1909

Brachyoxylon holbavicum sp. nov.

Fig. 5 a-i; Fig. 6 a-i.

The studied material

The studied material is represented by 24 decimetric-sized samples of silicified wood found around Holbav locality area (South Braşov County), on Maiului brook, in the Early Jurassic continental pyroclastic deposits considered of Pliensbachian age. All the studied samples are trunk or thick-branch fragments, with decimetric size, dark to black color and by magnifying glass or even by naked eye, regular fibrous structure without vessels suggesting a conifer wood. The specimens with field numbers: 1006, 1007, 1008, 1009, 1010, 1012, 1013, 1014, 1021, 1022, 1024, 1025, 1026, 1028, 1030, 1031, 1033, 1034, 1035, 1039, 1044, 1045, 1046, 1047 - belong to "Grădinaru Collection", are deposited now in the collections of the Geological Institute of Romania (G.I.R. Collections), under the inventory numbers 27612 - 27635 respectively, at the National Geological Museum, in Bucharest.

Microscopic description

In a synthetic description, all the studied specimens present secondary wood with tracheidoxylic structure devoid of normal resin canals, with a slight gradual diminishing of tracheids' size from the early to the late wood (very obvious in the specimens 1012, 1033), with relatively less distinct up to clearly distinct growth rings boundaries, marked by few rows of smaller flattened tracheids of the late wood. However in a specimen (i.e. 1010), an isolated canal rounded by a parenchyma cells was observed in cross section, most probably of traumatic origin and keeping inside some granular dark content. In some other specimens the thinner walled early and transitional wood is slightly compressed, or crushed or even collapsed, like in the charcoal (in the specimens: 1012, 1013, 1014, 1026, 1028). Sometimes, few details can be seen due to bad preservation during fossilization processes leading to vitrinite (as in the specimen 1029 which was excluded from study since it has no any structure kept, or only partially in the specimen 1031). Only in the sample 1044 appear the axial zone with primary wood that in cross section shows also arthropods galleries with coprolites, visible especially in the longitudinal sections. Tracheids with polygonal, mainly quadrangular cross section with rounded corners are displaying quadrangular slightly rounded lumina, or with curled walls often giving a "star-like" aspect of the lumina, often with discontinued walls due to the fossilization processes that often lead to deformed lumina or all of them are crushed by lateral compression (like in the specimen 1047). In a normal structure the tracheids, unequal in size, i.e. gradually diminishing from the early to the late wood, with radial/tangential diameters of 25-50 (60)/20-

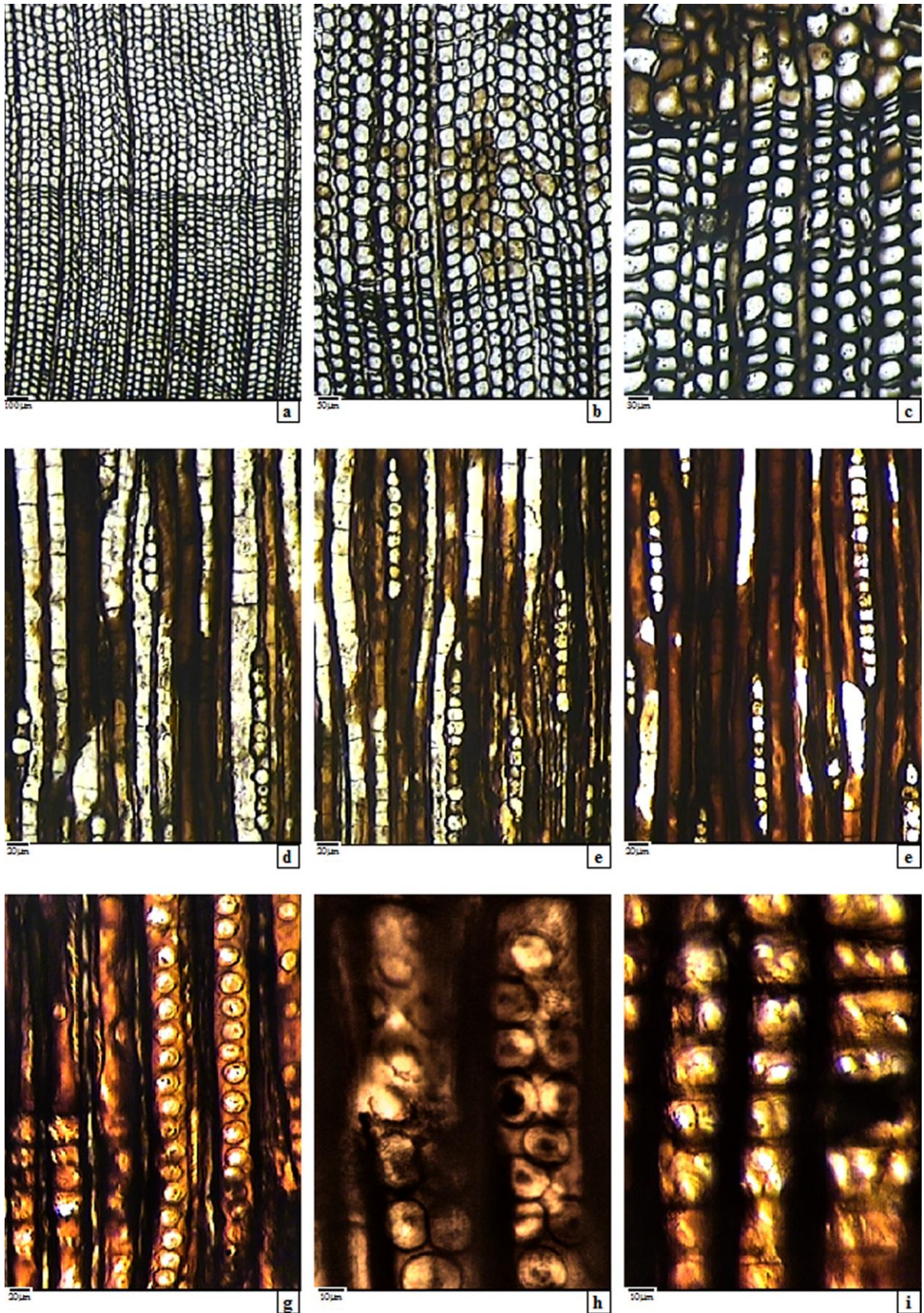


Fig. 5 *Brachyoxylon holbavicum* sp. nov. - holotype - inventory no. 27614 (field no. 1008). **a-c** Cross-section: distinct growth-ring boundary, thick-walled tracheids polygonal with rounded corners; linear rays. **d-f** Tangential section: unpitted tracheids, parenchyma, uniseriate rays. **g-i** Radial section: radial pitting of mixed type on tracheids, araucarioid cross fields with 2-3 superposed rows of cupressoid pits.

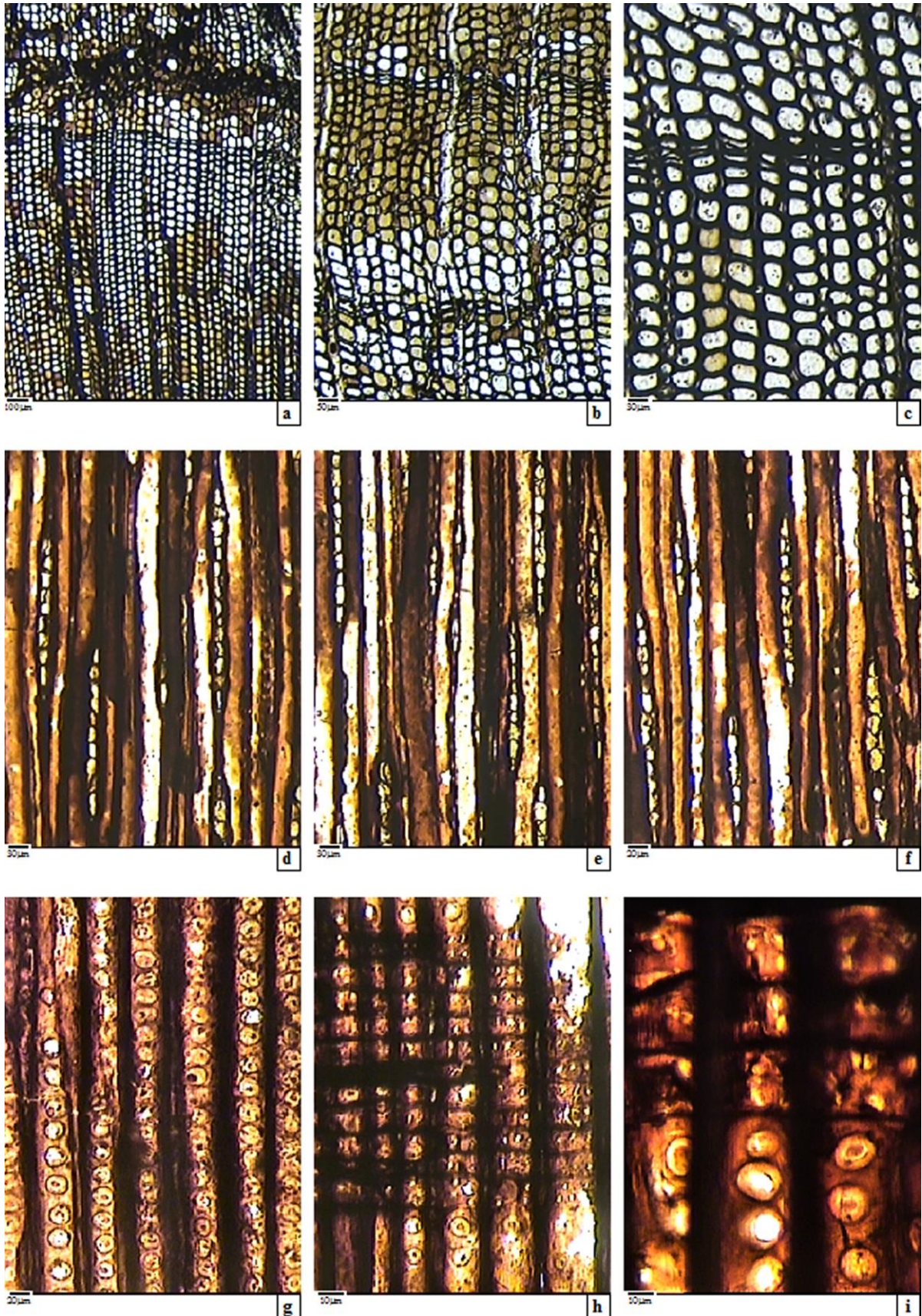


Fig. 6 *Brachyoxylon holbavicum* sp. nov. - paratype - inventory no. 27635 (field no. 1039). **a-c** Cross-section: quite distinct growth-ring boundary, thick-walled tracheids polygonal with rounded corners, and linear rays. **d-f** Tangential section: unpitted tracheids, uniseriate rays with biseriate storeys. **g-i** Radial section: radial pitting on tracheids of mixed type, araucarioid cross fields with 2-3 superposed rows of cupressoid pits.

40(60) μm and with thinner walls (3.5)5-7 μm (double wall) in the early wood and thicker in transitional wood, of 7-9 μm , to very thick in the late wood, of 10-13 μm the double wall. Between two successive rays there are bundles usually of 1-12 radial regular rows of normal tracheids, sometimes with obvious inter-cellular spaces (in the specimen 1008, 1034), and the density is of (550)836-1520 tracheids per square millimeter, even if sometimes is difficult to appreciate their number or to count them. Tangentially the tracheidal walls are usually unpitted, only sometimes seem to be uniseriately pitted with round or slightly flattened pits (smaller than on the radial walls), of 14-17 μm in diameters, contiguous or slightly spaced (in the specimens: 1026, 1030, 1047). The radial pitting is of mixed type, both uniseriate, with round pits, spaced or contiguous in a single vertical row and variably flattened (flattening index: $d/D=0.56-0.79-0.88$), or, (less) biseriately, round to oval, opposite to/or alternate, sometimes contiguous, combined with short uniseriate portions (in 1012). Generally, the pitted areas don't cover integrally the tracheidal wall. Also, when biseriately, round to oval pits appear, opposite or alternate, sometimes contiguous, combined with short uniseriate portions (in the specimens 1008, 1010). The pits have round to slightly oval borders of 19-24 μm in diameter to horizontal elliptic shape (when flattened), of 15.5-22 μm in diameters, and with round to oblique-elliptic apertures of 7-10(2)3-5 μm . There are no crassulae or helical thickenings. In the specimen 1034 the tracheids seems to show striations on the walls. Rarely (in specimen 1010), a solitary isolated canal was seen, rounded by epithelial parenchyma cells, probably of unknown traumatic origin, situation observed only in cross section. Inside the canal, some granular dark content appear.

A x i a l p a r e n c h y m a - is absent.

M e d u l l a r y r a y s - in cross section seen are thin, have a linear trajectory and are constituted from rectangular cells radially elongated, with smooth unpitted horizontal walls. Tangentially the rays are exclusively uniseriate, having 1-25 cells in height, or more, i.e. 25-600(650) μm . Sometimes the taller rays have 1-2 short biseriately storeys of the same thickness that one cell, more numerous in specimen 1046. The ray-cells are polygonal-rounded to oval (in height), usually having relatively not too thick walls and, sometimes, obvious lateral empty spaces (in specimens 1007, 1013, 1014, 1024, 1026, 1031, 1034). The ray-density is of 7-12 rays per tangential horizontal millimeter. In radial section the rays are homocellular, constituted by cells all procumbent of 20-24 μm in height. In the marginal rows, the cells are taller, of 28-35(-50) μm . The ray cells have moderately thick horizontal walls, of 4-6.2 μm the double wall, sometimes more, and the outer wall of the marginals is slightly curled. The tangential wall is thick and there are no visible indentures. The typical araucarioid cross-fields have 1-6 oculipores of cupressoid, tending to podocarpoid, usually hexagonal-rounded to oval, with diameters of of 9-15(18) / 6-8(10) μm , with tilted elliptic apertures of 5-8 / 3-5 μm . The pits arrangement is usually alternate as 1-6(9) pits on 1-2(3) horizontal rows in the body ray cells and 2-9 pits on 2-3 horizontal rows, alternate or slightly irregular, in the marginal fields

Affinities and discussions

All the here studied 24 samples of fossil wood present tracheidoxyl structure and mostly uniseriate rays, pitting of mixed type on the radial tracheidal walls and araucarioid cross-fields, xylotomical details that send to *Brachyoxylon* Hollick et Jeffrey, a type of Mesozoic tracheidoxyl.

The type species for this genus, *Brachyoxylon notabile* Hollick et Jeffrey, 1909 has also radial tracheidal pitting of mixed type and araucarioid cross-fields. The presence of traumatic resin canals is considered as a result of injury (see table 1). Otherwise, in a synthesis of Taylor & Taylor (1993, p. 689) is noted that no "cheirolepidiaceus" wood has normal resin canals but traumatic canals can be present sometimes in this genus, as a result of injury or freezing (see also Philippe, 1995).

So, the description of our studied structures agrees with the diagnosis of the genus *Brachyoxylon* Hollick et Jeffrey, even if the original diagnosis, after the designated generotype (*B. notabile* Hollick et Jeffrey, 1909) is very elliptic, defining "a tracheidoxyl devoid of normal secretory ducts, with mixed radial pitting, araucarioid cross-fields and other ray walls integer" (see Philippe, 1993; Philippe & Bamford, 2008).

To do a comparison with fossil forms already described and published, it is important to remind that some species are no more considered as *Brachyoxylon* species, so must be excluded from a discussion, as for example:

- *Brachyoxylon nipponicum* Nishida, that according Vozenin-Serra & Pons (1990) is synonymous with *Protopodocarpoxylon orientale* Serra, 1969 (see also Machhour & Pons, 1992);

- *Brachyoxylon brachyphylloides* (Torrey) Kräusel, according Philippe (1993) is in fact a species of *Telephragmoxylon*: *T. brachyphylloides* (Torrey) Philippe;

- *Brachyoxylon notabile* Hollick et Jeffrey, 1909, *B. woodworthianum* Torrey, 1923, *B. libermanii* Philippe, 1995 preserve also primary structure and pith and such structures were described as *Telephragmoxylon* (by Torrey, 1921) or as *Brachyoxylon* sp. (by Holden, 1913a; Shimakura, 1937), having mostly uniseriate radial pitting and low rays (according to Machhour & Pons, 1992);

- *Brachyphyllum desnoyersii* as initially was described by Lemoigne (1968) was renamed as *Agathoxylon desnoyersii* (Lemoigne) Philippe, 1995, and now considered as a junior taxonomical synonym (Philippe *et al.*, 2018).

- Also, *Brachyoxylon urkutense* Greguss according Philippe & Barbacka (1997) it is synonymous with *Simplicioxylon hungaricum* Andreánszky, 1949.

Two older described taxa from North America as *Paracedroxylon scituatense* Sinnot and *Voltzioxylon dokumense* Torrey were sent into synonymy with *Brachyoxylon*, as *B. scituatense* (Sinnot) Philippe, 1993 and *B. dokumense* (Torrey) Philippe, 1993 - respectively, having typical characters of this genus.

Anyway, we tried to describe more accurately the very specific features of this big group of specimens (24) which could define a new species. All these characters were introduced into a comparative table of other described species of *Brachyoxylon* (Table 1), taken from Bodnar *et al.* (2013), and we observed that our described material is slightly different, by the rare apparition of trau-

Table 1 Comparative table of species of the genus *Brachyoxylon* (after Bodnar *et al.*, 2013).

| Species of <i>Brachyoxylon</i> | Age | Distribution | False growth rings | Resin canals | Tangential tracheid pitting | Radial tracheid pitting | Axial parenchyma | Cras-sulae | Rays (height) | Cross-field pits |
|--|----------------------------|-----------------------------|--------------------|-----------------------------|-----------------------------|--|---------------------|------------|--|--|
| <i>B. notabile</i> Hollick et Jeffrey, 1909 | Upper Cretaceous | USA | absent | traumatic | present | mixed 1–2 seriate | absent | present | uniseriate (1–8 cells) | 5–11 cupressoid |
| <i>B. woodworthianum</i> Torrey, 1923 | Lower Cretaceous | USA | absent | absent | absent | mixed | absent | absent | uniseriate (high) bi-pluri-seriate(low) | cupressoid |
| <i>B. brachyphylloides</i> (Torrey, 1923) Kräusel, 1949 | Middle Jurassic-Cretaceous | Tunisia, Nigeria, Chad, USA | absent | traumatic | present | mixed 1–2 seriate | absent | absent | uniseriate, partially biseriate (1–20 cells) | 1–10 cupressoid |
| <i>B. japonicum</i> (Shimakura, 1937) Kräusel, 1949 | Lower Cretaceous | Japan | absent | absent | present | mixed 2–3 seriate | absent | absent | uniseriate (50–60 cells) | cupressoid |
| <i>B. saurinii</i> Boureau et Serra, 1961 | Jurassic (?) | Cambodia | absent | traumatic | absent | mixed 1–2 seriate | absent | absent | uniseriate (1–31 cells) | 2–12 cupressoid |
| <i>B. boureaui</i> Serra, 1966 | Jurassic (?) | Cambodia | absent | absent | absent | mixed 1–2seriate | absent | absent | uniseriate (low) | 4–36 cupressoid |
| <i>B. nonakai</i> Yamazaki et Tsunada, 1981 | Upper Jurassic | Japan | absent | absent | absent | opposite 2–4 seriate | absent | absent | uniseriate (unknown) | cupressoid |
| <i>B. lagonense</i> (Laudouén.) Dupéron-Laudouéniex 1991 | Cretaceous | Chad | absent | traumatic | absent | mixed 2–3 seriate | absent | absent | uniseriate (low) | cupressoid |
| <i>B. trautii</i> (Barale, 1981) Philippe, 1995 | Middle Jurassic | France | present | absent | present | mixed 1–2 seriate | absent | absent | uniseriate (low) | 4–9 cupressoid |
| <i>B. liebermanii</i> Philippe, 1995 | Jurassic | France | present | absent | present | mixed 1 seriate | absent | absent | uniseriate (low) | 5–12 cupressoid to podocarpoid |
| <i>B. eboracense</i> (Holden, 1913a) Philippe, 2002 | Jurassic | England | absent | absent | absent | mixed 1–2 seriate | absent | absent | uniseriate (low) | numerous cupressoid |
| <i>B. orientale</i> (Serra, 1969) Philippe <i>et al.</i> , 2004 | Lower Cretaceous | Tibet | present | absent | absent | mixed 1–2 seriate | absent | absent | uniseriate (low) | 3–10 podocarpoid |
| <i>Brachyoxylon</i> sp. A Philippe <i>et al.</i> , 2004 | Late Jurassic | Thailand | present | absent | present | mixed 1–2 seriate | absent | absent | uniseriate (low) | up to 12 cupressoid |
| <i>B. semibiseriatum</i> (Pant et Singh, 1987) Kurzawe et Merlotti, 2010 | Permian | Brazil | absent | present | present | mixed 1–5 seriate | absent | absent | uniseriate (1–38 cells) | 4–16 cupressoid |
| <i>B. serrae</i> Phillippe <i>et al.</i> , 2011 | Lower Cretaceous | Thailand | absent | traumatic | present | mixed 1–2 seriate | absent | absent | uniseriate (1–15 cells) | 5–16 cupressoid |
| <i>B. avramii</i> Iamandei et Iamandei, 2005 | Early Cretaceous | Romania | absent | traumatic | rare | mixed, 1–2 seriate | absent | absent | uniseriate (1–21 cells) | 1–8 cupressoid to podocarpoid |
| <i>B. dobrogiacum</i> Iamandei et Iamandei, 2005 | Early Cretaceous | Romania | absent | absent | present | mixed 1–3 seriate | present | present | uniseriate (1–10 cells) | 1–6 cupressoid to podocarpoid |
| <i>B. holbavicum</i> sp. nov. 24 specimens | Lower Jurassic | Romania | absent | absent (+ rarely traumatic) | absent, or rarely present | mixed type (1-seriate + partially 2-seriate) | absent | absent | uniseriate + biseriate storeys [1–25 cells tall] | 1–9 pits cupressoid to podocarpoid + alternate arrangement |
| <i>B. cristianicum</i> sp. nov. 8 specimens | Lower Jurassic | Romania | absent | absent | absent | mixed 1–2 seriate | present (+crystals) | absent | uniseriate + partially biseriate [1–20 cells tall] | 1–6 pits cupressoid to podocarpoid +alternately arranged |

matic canals, the absence of the axial parenchyma, and the presence of the exclusively uniseriate rays, relatively tall, of up to 25 cells.

There are numerous already described forms of *Brachyoxylon*, from different regions of the Earth, with various very specific details. However, we observed that *B. saurinii* of Boureau & Serra (1961), *B. avramii* and *B. dobrogiacum* of Iamandei & Iamandei (2005) and *B. serrae* of Phillippe *et al.* (2011) have high rays of more than 20 cells and uniseriate radial pitting, quite similar to our studied material. The both species described by Iamandei & Iamandei (2005) from Romania are branch fragments preserving also primary wood, like the specimen 1044 studied here which also shows arthropods galleries with coprolites, but the characters of the secondary wood are not similar.

Many other forms of *Brachyoxylon* were described in the world, showing a large spread of this Mesozoic genus on the planet, and we only quote some of them:

- Iijima *et al.* (1989) described a species of *Brachyoxylon*, from Far East, from the uppermost Triassic - Lower Jurassic, in the Aoyama Chert from Kuzuh, Japan, by the study of a drifted piece of wood with this type of structure.

- Cevallos-Ferriz (1992) described some other species of *Brachyoxylon* from the Upper Cretaceous from north of Mexic.

- Philippe *et al.* (2004a) cited above, made a complex biogeographic analysis of Jurassic–early Cretaceous wood assemblages, *Brachyoxylon* included, from Gondwana.

- Hickey *et al.* (2011) described a typical *Brachyoxylon* from the Upper Triassic from Connecticut, USA.

- Also, Vera & Césari (2012) have described five specimens of Aptian fossil woods from Argentina, identified as *Brachyoxylon* sp. cf. *Brachyoxylon boureaui* Serra, 1966 characterized by radial mixed pitting as predominantly uniseriate rows of circular to polygonal bordered pits. The biseriate and triseriate pitting pattern is also

present, but is less common. The cross-field regions show 8-26 small circular pits arranged in 2-4 rows, with slit-like pit apertures, obliquely oriented, in rays exclusively uniseriate, and relatively low.

- Bodnar *et al.* (2013) described a *Brachyoxylon currumilii* from the lower-middle Jurassic of Argentina and made a comparative table of already published species of *Brachyoxylon* (Table 1), in which we added the xylotomical characters of our here studied material, for the comparison of the essential xylotomical features.

- Garcia *et al.* (1998) made a debate on the Jurassic palaeoenvironment of large European forests of *Agathoxylon* and *Brachyoxylon* described on the France territory.

- More recently, Philippe *et al.* (2018) reidentified from "Lignier collection" a sample initially identified by Lignier as *Cedroxylon blevilense*, as new combination: *Brachyoxylon blevilense* (Lignier) Philippe *et al.*, 2018 - based on the mixed type of radial pitting and the araucarioid cross-fields. Also they identified the "Échantillon n°150" as having typical araucarioid cross-fields, with up to six crowded oculipores and resiniferous axial parenchyma (traumatic?), similar to *Brachyoxylon liebermannii* Philippe, a species previously described by Philippe (1995), from the Liassic of northeastern France.

So, considering the xylotomic matches of our studied material and the already published species from the comparative table (see Table 1) and the comparison with the original description of other species, we concluded that our material (which presents a secondary wood structure usually devoid of resin canals, even if rarely of traumatic type can appear, with mixed radial pitting, with axial parenchyma usually, with uniseriate rays or partially biseriate, relatively high, up to 25 cells and and cross-fields with oculipores of cupressoid type, tending to podocarpoid, in an araucarioid arrangement) has a structure of *Brachyoxylon* of a special type.

Here, in our paper, other very similar form of *Brachyoxylon* was described (i.e. *B. cristianicum*), slightly similar and coming almost from the same areal (see Table 1). So, taking into account these observations and discussions, we consider we have described here a new species of *Brachyoxylon*, that we named *Brachyoxylon holbavicum* sp. nov., after the name of the locality area (Holbav) from where it was found.

From the studied material it was designated as holotype the specimen with inventory number: 27614 (field no. 1008) and a paratype with inventory number: 27635 (field no. 1039) from a population of 24 specimens of fossil wood from "Grădinaru Collection" with similar xylotomic characters to almost identical. The mentioned types and all the studied material are now deposited in G.I.R. Collection, in the National Geological Museum in Bucharest,

Diagnosis of the new species *Brachyoxylon holbavicum* sp. nov.:

The secondary wood of tracheidoxylic type has distinct growth rings, sometimes with rare, isolated, traumatic canals, thick-walled tracheids, polygonal in cross-section gradually diminishing in size to the late-wood. Tangentially the tracheids are usually unpitted, rarely with smaller pits contiguous or slightly spaced. The radial pitting is of mixed type, usually uniseriate with round pits, spaced or contiguous and variably flattened and, less, biseriate, with

round to oval, opposite to alternate, sometimes contiguous pits, combined with short uniseriate portions. Pits with round to oblique-elliptic apertures. Axial parenchyma is absent. Rays exclusively are uniseriate, tall, with biseriate storeys. Rays exclusively uniseriate, tall, with biseriate storeys, with lateral empty spaces and relatively thick walls. Homocellular, with curled outer wall, and without indentures. Cross-fields typical araucarioid, 1-6(9) oculipores hexagonal-rounded to oval of cupressoid type tending to podocarpoid in 1-3 horizontal rows arranged, alternate to slightly irregular, in the marginal fields on 2-3 horizontal rows.

Brachyoxylon cristianicum sp. nov.

Fig. 7 a-i; Fig. 8 a-i.

The studied material

The studied material is represented by 8 pieces of silicified wood, four of them found in the Cristian area, on the Fabricii valley, and four in the Holbav area, on the Maiului brook, in the Lower Jurassic continental deposits with pyroclastic character considered of Sinemurian-Pliensbachian age, most probably Pliensbachian. All the studied samples are trunk or thick-branch fragments, having decimetric size, dark to black color and, by magnifying glass or even by naked eye, regular fibrous structure without vessels can be seen, clearly suggesting a conifer wood. The specimens having field-numbers: 1015, 1016Cr, 1018Cr, 1019Cr, 1020Cr, 1027, 1042, 1043 belong to "Grădinaru Collection", and are deposited now in the collections of the Geological Institute of Romania (G.I.R. Collections), under the inventory numbers: 27636 –27643 respectively, at the National Geological Museum, in Bucharest.

Microscopic description

All the eight studied specimens present secondary wood with tracheidoxylic structure with quite distinct growth-rings boundary marked by the late wood which has smaller tracheids than the early wood. The structure is devoid of any normal or traumatic resin canals.

Tracheids - in cross section viewed are polygonal with slightly rounded corners, sometimes determining intercellular spaces (specimens 1020, 1027), unequal in size, with quadrangular to rounded large lumina with radial/tangential diameters of 20-42(50-60)/20-35-(45-65) μm (smaller in the late wood), and with relatively not too thick walls: (4)6-9 μm the double wall in the early and in the transitional wood, up to 9-13 μm the double wall in the late wood. The interradian bundles have 1-10(12) radial regular rows of tracheids and density is 992-1444(1518) tracheids per mm^2 . Tangentially seen the tracheids are usually unpitted, or rarely pitted (in specimen 1019) with small round pits ($d=8-12 \mu\text{m}$) with small apertures ($d=4-5 \mu\text{m}$), in a single row arranged, slightly irregular, locally as two opposite pits. The radial pitting is of mixed type, with almost abietinean pits, arranged either in a single vertical row as round pits, spaced, or contiguous separated by a horizontal line, with round or slightly elliptic apertures, or biseriate, round, alternate or opposite, contiguous on short rows often continued by uniseriate portions, not occupying all the length of the wall. In fact, generally, the pitted areas don't cover integrally the tracheidal wall.

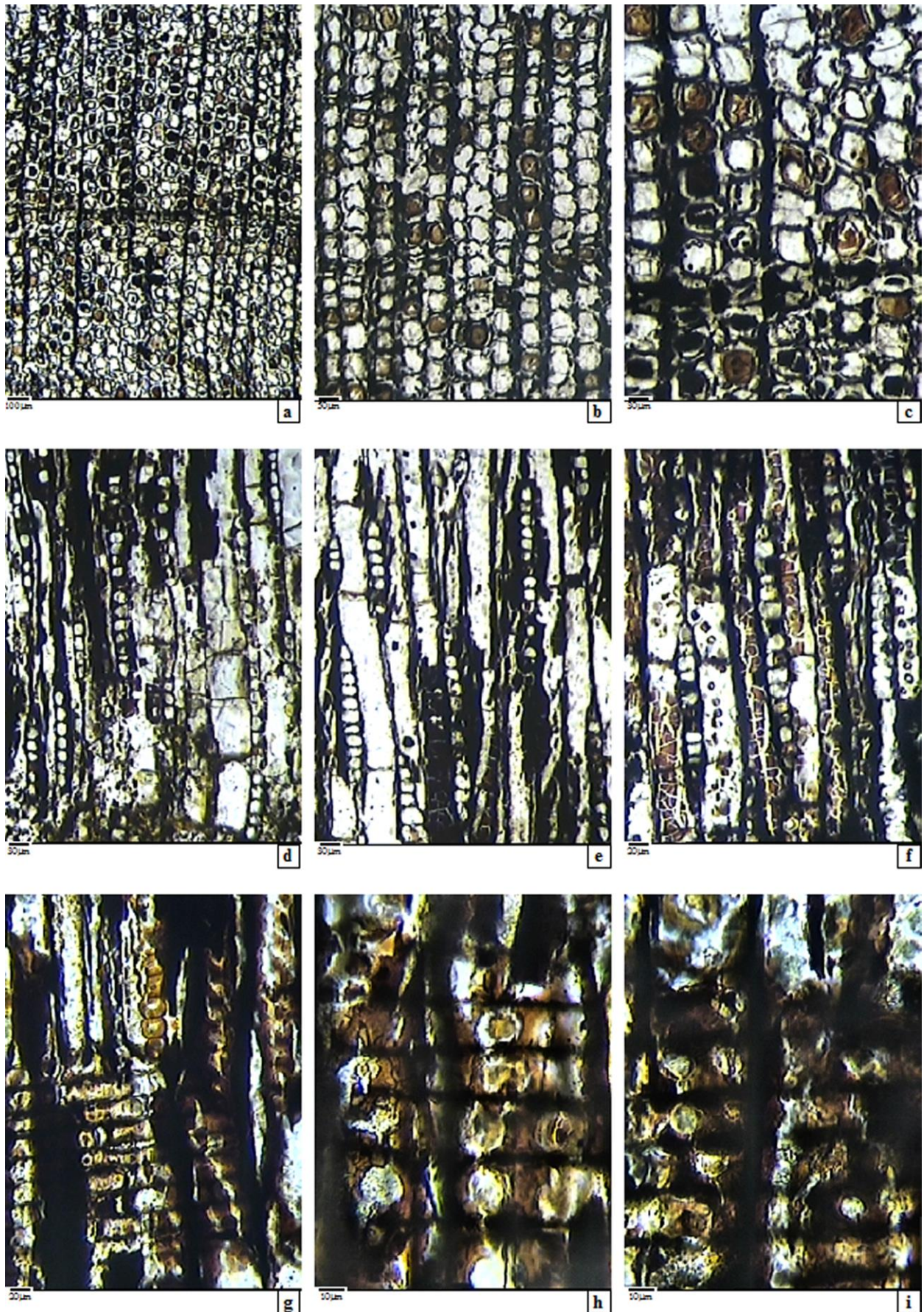


Fig. 7 *Brachyoxylon cristianicum* sp. nov. - holotype - inventory no. 27639 (field no. 1019cr). **a-c** Cross-section: distinct growth ring boundary, thick-walled tracheids, polygonal with rounded corners, and linear rays. **d-f** Tangential section: unpitted tracheids, parenchyma with pits and crystals; uniseriate rays. **g-i** Radial section: uniseriate radial pitting on tracheids of mixed type, cross fields with 1-2 rows of oculipores.

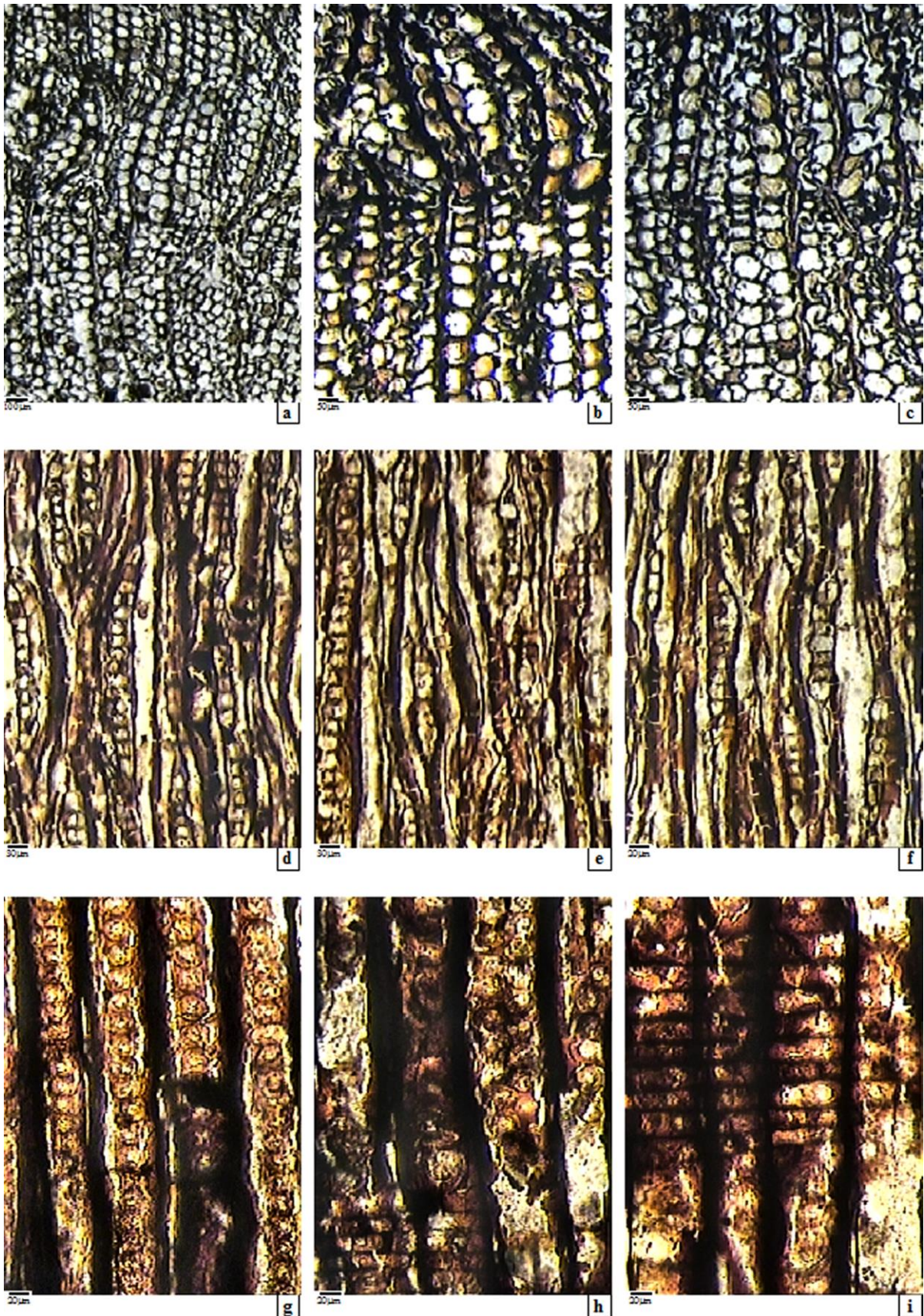


Fig. 8 *Brachyoxylon cristianicum* sp. nov. - paratype - inventory no. 27640 (field no. 1020cr). **a-c** Cross-section: distinct growth-ring boundary, polygonal tracheids with rounded corners and thick walls; linear rays. **d-f** Tangential section: unpitted tracheids, parenchyma with pits and uniseriate rays. **g-i** Radial section: radial pitting on tracheids of mixed type, uniseriate and biseriate, cross fields with 1-2 rows of oculipores.

The pits have round borders of 18-22 often flattened when contiguous ($d/D=0.65-0.88$), and round or tilted to lying elliptic apertures of 7.5-10/2-3 μm (in specimen 1019). None crassulae or helical thickenings have been seen, but sometimes a brown content is present.

A x i a l p a r e n c h y m a - in cross-section view is present, few, dispersed, often difficult to identify amongst the tracheids. It has relatively thin vertical walls relatively similar to those of the tracheids and sometimes with simple pits, irregularly arranged on 1-2 rows. The horizontal (terminal) walls are thin and smooth or slightly nodular and, inside the cells, resin content or plugs or granules, and sometimes small polygonal crystals are present (specimens 1015, 1019).

M e d u l l a r y r a y s - in cross section appear usually uniseriate, rectilinear, and with smooth horizontal cellular walls. In tangential view the rays have 1-20 (sometimes up to 25) cells in height, i.e. 30-570(650) μm high. The taller uniseriate rays have biseriations, sometimes giving aspect of fusiform rays (specimens 1015, 1016). The ray cells are polygonal rounded to oval sometimes having dark content and relatively thin walls. Ray density is 5-11 rays per horizontal tangential millimeter. Radially the rays are homocellular, cells all procumbent of 20-30 μm tall, moderately thin walled, of 4-5(7) μm the double wall. Within the marginal rows the cells are taller, of 25-35(40) μm , and have corrugated outer wall. The cross-fields of araucarioid type have 1-6 oculipores, hexagonal or slightly rounded to oval, of cupressoid type, tending to podocarpoid, with 15-19/8-12 μm in diameters, having circular or short elliptic tilted apertures of 10-14/5-7 μm . The pits arrangement is in 1-2 rows in the ray body fields, in the marginal fields on two horizontal rows, always alternate to slightly irregular when on two rows.

Affinities and discussions

After the microscopic study of each sample and the synthetic evaluation of their generic characters, a number of 8 samples of fossil wood were grouped all of them having tracheidoxyl structure, uniseriate rays, radial pitting of mixed type on the tracheids and araucarioid cross-fields, suggesting a *Brachyoxylon* type of wood slightly different of the above described form (see Table 1).

The discussion on the status of the genus *Brachyoxylon* Hollick et Jeffrey, and of the species described under it along the time is presented in detail above. Even if the type species (*Brachyoxylon notabile* Hollick et Jeffrey, 1909) has a very elliptic diagnosis, referring only to "a tracheidoxyl devoid of normal secretory ducts, with mixed radial pitting, araucarioid cross-fields and other ray walls integer" (see Philippe, 1993), it is interesting that traumatic canals were described, considered as a result of injury or freezing (Taylor & Taylor, 1993, p. 689; Philippe, 1995). Anyway, the specimens described here are devoid of any kind of canal.

In our studied specimens, apart from the known characters of the genus, there are some specific xylotomic details regarding the size of tracheids or their behavior inside the growth ring, their pitting type, aspect and size of the pits on the tracheids and in the cross fields. The accurate description of these very specific features of the stu-

died material can be useful for the comparison with the already published and valid species (see Table 1).

Anyway, it seems that the presence or the absence of the axial parenchyma and of the crassulae, and also the ray heights and the pits' number in the cross-fields, are important to separate the species of *Brachyoxylon* (Philippe, 1993, 1995).

- Two forms of *Brachyoxylon* preserving primary structure and pith were described from the Lower Cretaceous in southeastern of Romania as *B. avramii* and *B. dobrogiacum* (Iamandei & Iamandei, 2005) as having a tracheidoxyl structure with few parenchyma and uniseriate low rays, with mixed radial pitting and araucarioid cross-fields (sensu Philippe, 1993), bearing cupressoid to podocarpoid pits in an araucarioid arrangement, but are slightly different of the here studied material which has high rays with biseriate stories, even with aspect of biseriate fusiform rays.

- Also, Vera & Césari (2012) have described an interesting material represented by five specimens of Aptian fossil woods from Argentina and identified as *Brachyoxylon* sp. cf. *Brachyoxylon boureaui* Serra, 1966 having radial pitting of a mixed type, predominantly uniseriate, with circular to polygonal bordered pits, sometimes biseriate and triseriate, even if less common, and numerous pits in the cross-fields (8 up to 26 small circular pits arranged in 2-4 rows, with slit-like pit apertures, obliquely oriented), details which are very different of our specimens.

- The Jurassic new species of *Brachyoxylon currumilii* described by Bodnar *et al.* (2013) from Argentina, has also some details which differ of our material. But in that paper there is a very well documented discussion on the status of *Brachyoxylon* genus and its already described species, and a comparative table of the valid species of *Brachyoxylon* in which we added, for comparison, the essential xylotomic characters of our studied material (Table 1).

- Interesting for comparison seems to us the more recently reidentification of Philippe *et al.* (2018) from "Lignier collection" of a sample initially identified as *Cedroxylon blevilense* by Lignier, as *Brachyoxylon blevilense* (Lignier) Philippe *et al.*, 2018 based on the mixed type of radial pitting and the araucarioid cross-fields. Also the identification of the "Échantillon n°150" from the same collection, which showed typical araucarioid cross-fields with up to six crowded oculipores and traumatic(?) resiniferous axial parenchyma is similar to *Brachyoxylon liebermannii* Philippe, 1995, a species previously described from the Liassic of northeastern France.

But, considering the xylotomic similarities and dissimilarities of our studied material with the already published species from this comparative table (Table 1) and the comparison with the original description of the species, and also, the discussions from above, it can be observed that our material is special.

Thus, it presents a secondary wood structure without resin canals, with mixed radial pitting, uniseriate spaced or contiguous, or biseriate, alternate or opposite, contiguous, with round to elliptic apertures, with axial parenchyma, with uniseriate rays or even biseriate, relatively high, and

cross fields with cupressoid oculipores in an araucarioid arrangement.

These xylotomical details are slightly different of the other species of *Brachyoxylon* described in this paper and, based on this difference well expressed in the Table 1 from above, we considered we have described here a new species that we named *Brachyoxylon cristianicum* sp. nov., after the name of the region from where it was found (Romanian spelling for Christian is Cristian). From the studied material a holotype was designated the specimen with field number 1019 in "Grădinaru Collection", (inventory number 27639), and a paratype the specimen with field number 1020 in "Grădinaru Collection" (inventory number 27640), from a population of 8 specimens of fossil wood with similar to almost identical xylotomic, all material now deposited in GIR Collection, in the National Geological Museum, Bucharest, as we specified above.

Diagnosis of the new species *Brachyoxylon cristianicum* sp. nov.:

Tracheidoxyl with distinct growth rings, devoid of any canal. Tangentially the tracheids are usually unpitted, rarely with small pits, uniseriate. Radial pitting of mixed type, of pits abietean but with araucarian arrangement, uniseriate, spaced or contiguous, with round or slightly elliptic apertures, or biseriate, alternate or opposite, contiguous and flattened, in short rows often continued by uniseriate portions, not occupying all the length of the wall. Axial parenchyma present, dispersed, in longitudinal view presenting thin and smooth or slightly nodular horizontal walls and resin content or even crystals. Rays uniseriate, relatively tall, with biseriations, sometimes with fusiform aspect, homocellular, having typical araucarioid cross-fields with 1-6 oculipores alternate to slightly irregular on 1-2 rows, hexagonal rounded to oval, of cupressoid type tending to podocarpoid, with round or short elliptic tilted apertures.

Genus *Protophyllocladoxylon* (Kräusel, 1939) Vogelgehner, 1968

Protophyllocladoxylon holbavicum sp. nov.

Fig. 9 a-i.

The studied material

The studied material is represented by a single decimetric sized sample of silicified wood found in Holbav area, on the Maiului brook, in Lower Jurassic continental pyroclastic deposits considered of Pliensbachian age. The studied sample represents a trunk fragment, having dark to black color and by magnifying glass or even by naked eye, regular fibrous structure without vessels is visible, suggesting a coniferous wood. The specimen with the field number 1040 belong to "Grădinaru Collection" and is deposited now under the inventory number 27644 in the collections of the Geological Institute of Romania (G.I.R. Collections), hosted by the National Geological Museum, Bucharest.

Microscopic description

The studied specimen has secondary wood with tracheidoxylic structure devoid of any normal or traumatic resin canals, with diminishing tracheid size only in the last part of the transitional wood close to the late wood,

which is represented by few rows of smaller tracheids, marking distinctly the growth ring boundary.

T r a c h e i d s - with polygonal in cross-section and polygonal lumina, slightly rounded at corners, have radial/tangential diameters of 20-40/20-45 μm and relatively thick walls, of 4-10 μm double wall. The interradian bundles have 1-9 radial regular rows of tracheids and the density is 784-899 tracheids per mm^2 . Tangentially seen the tracheids are usually unpitted, but sometimes small circular pits of 8-12 μm in diameter, in a single vertical row appear. Radial pitting is of mixed type, almost exclusively uniseriate, with pits circular, spaced or contiguous, with round or oblique elliptic apertures. When contiguous they are slightly flattened. The pits have round borders of 17-21 μm and rounded to tilted elliptical apertures of 5-7/3-4 μm . There are no crassulae or helical thickenings.

A x i a l p a r e n c h y m a - appears in cross-section as few, disperse cells relatively thick-walled, difficult to distinguish of tracheids, but visible in tangential view, having thick terminal (horizontal) walls.

M e d u l l a r y r a y s - are fine, rectilinear in cross-section, with smooth horizontal cellular walls. Tangentially the rays have 1-21 cells in height, or more, i.e. 30-550 μm high. The taller rays have high biseriations (up to ten storeys) giving to the rays a biseriate aspect with long uniseriate endings. The ray cells are polygonal rounded to oval and have relatively thick walls. Ray density is 8-12 rays on horizontal tangential millimeter. Radially the rays are homocellular, cells all procumbent of 20-30 μm high, moderately thick-walled: 5-6.2 μm the double wall. The tangential walls are thick-walled or slightly nodular. Indentures indistinct or absent. The typical cross-fields have 1-2 oopores of phyllocladoid type, more or less bordered, and inclined, of 13-17 μm in diameters, with large lens-like apertures of 6.5-8.5/3-5 μm .

Affinities and discussions

The here studied sample of fossil wood presents tracheidoxylic structure with uniseriate rays, pitting of mixed type on the radial tracheidal walls and phyllocladoid cross-fields, aspects which suggest similitude with members of the Family Podocarpaceae Endlicher, 1847, especially with the extant *Phyllocladus* Rich. ex Mirb., and send to the fossil correspondent *Protophyllocladoxylon* (Kräusel) Vogelgehner, 1968, a type of tracheidoxylic Mesozoic fossil wood having such characters. In fact the species described from the Cretaceous of Syria (*P. leuchsii*) and designated as type-species by Kräusel (1939), has some problems of origin, age and anatomy (Philippe *et al.*, 2003; Pujana, 2005), so a reappraisal of this genus is necessary.

The correct diagnosis of the genus *Protophyllocladoxylon* (Kräusel, 1939) as is emended by Vogelgehner (1968), is presented by Philippe (1995) like this: "Tracheidoxyl without secretory canals, with radial pitting of mixed type and cross-fields with 1-2 elliptic oopores to large oval, more or less oblique, horizontal ray-cells walls integer". Or, as it's specified by Philippe & Bamford (2008), "oopore more phyllocladoid (i.e. pointed and oblique), usually somewhat bordered in the latewood". In fact, large, oval, more or less oblique oopores are present in the cross-fields only to *Protophyllocladoxylon*, but not to

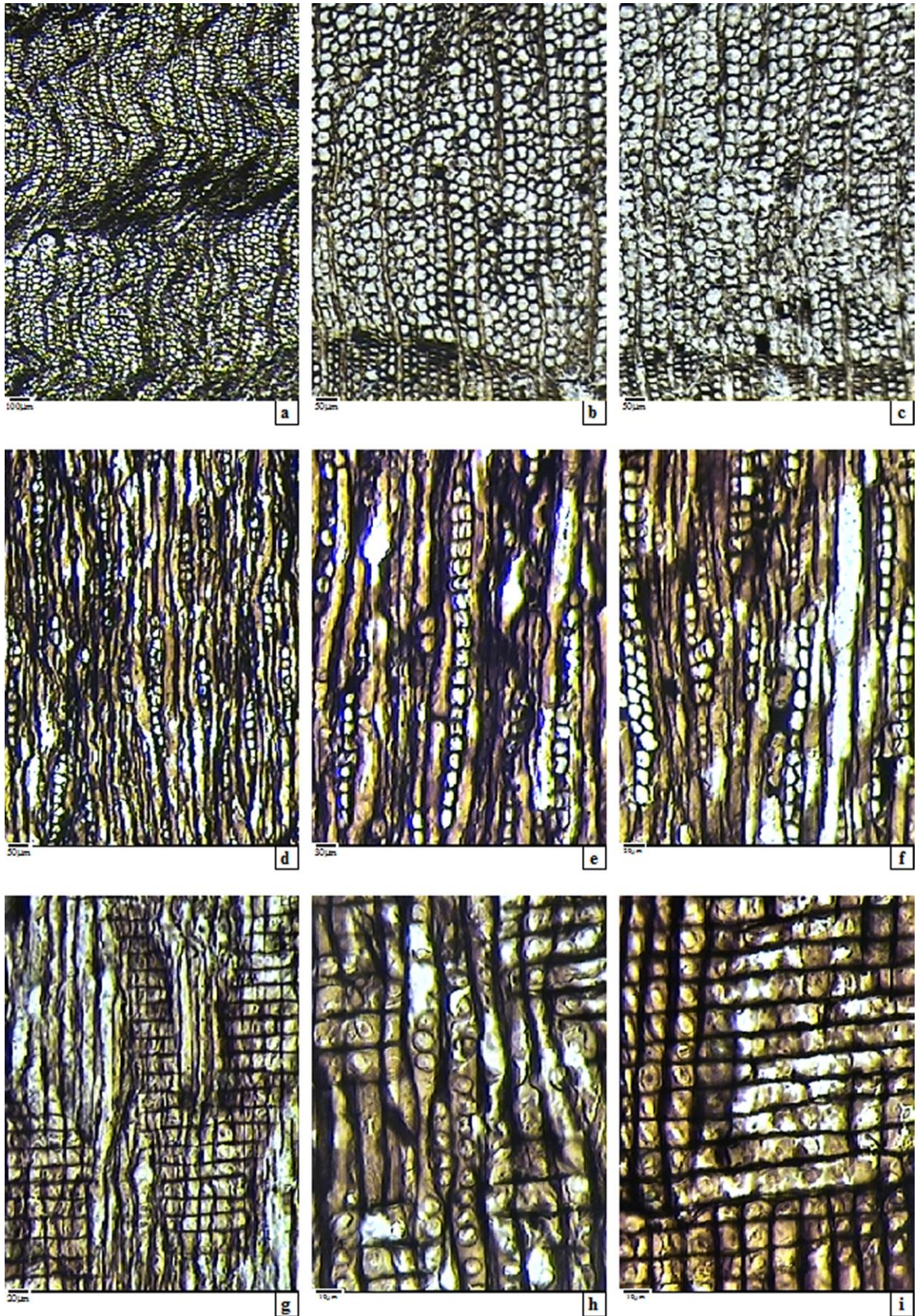


Fig. 9 *Protophyllocladoxylon holbavicum* sp. nov. - holotype - inventory no. 27644 (field no. 1040). **a-c** Cross-section: distinct growth-ring boundary, polygonal tracheids with rounded corners, and linear rays. **d-f** Tangential section: unpitted tracheids, uniseriate rays, and biseriate with long endings. **g-i** Radial section: uniseriate radial pitting on tracheids of mixed type, phyllocladoid cross-fields with 1-2 phyllocladoid large inclined oculipores.

Xenoxylon, *Protocircoporoxylo*n or *Prototaxodioxylo*n, where other types of oopores appear.

We tried a comparative analyse of our studied specimen with some already described forms from worldwide because during Paleozoic and Mesozoic times it seems that this genus had a widespreading, and we cite:

- *Paraphyllocladoxylon araucarioides* Holden, 1913b, was relatively badly described and, having araucarian characters, is possibly to be reconsidered as *Protophyllocladoxylon* (Philippe & Bamford, 2008). In fact, these authors specified that the groups of Paleozoic and Mesozoic described species of *Protophyllocladoxylon* have araucarian characters and must be recognized and a reappraisal of this genus is necessary.

- Also, numerous species of *Protophyllocladoxylon* described before 2003 felt in synonymy with *Metapodocarpoxylon libanoticum* (Edwards) Dupéron-Laudoueneix et Pons, aspect discussed in detail by Philippe *et al.* (2003) and Philippe & Bamford (2008).

- *Protophyllocladoxylon quedlinburgense* Schultze-Motel, 1961 has flattened radial uniseriate pitting and also biseriate opposite, so it's slightly different of our material.

- *Protophyllocladoxylon franconicum* was described from Germany by Vogellehner (1966) is a Mesozoic form with phyllocladoid type of oopores and with radial uniseriate pitting, rather similar to our specimen.

- More recently, Zhang Yi *et al.* (2010) have described a Paleozoic species: *Protophyllocladoxylon jingyuanense*, from Carboniferous (Late Mississippian) deposits from NW China, with similar features in radial pitting and in cross-fields.

- Boura *et al.* (2013) described some Early Cretaceous woods from Vietnam beside some species of other gymnosperms, a *Protophyllocladoxylon xalucense*, a new species bearing all the typical features of the genus and more than that, an abundance of axial parenchyma. Also, a documented comparison with numerous species of *Protophyllocladoxylon* described in far-eastern Asia was done, all of them different of our specimen.

- Also, Fletcher *et al.* 2014 have described a *Protophyllocladoxylon owensii* as a new species from the Upper Cretaceous of Queensland (Australia), having radial pitting on the tracheid predominantly alternate and contiguous (of araucarian type), uniseriate in latewood, biseriate and rarely triseriate, in earlywood, but the cross-field pits are typically phyllocladoid and have oopores or more rarely circopores, two per cross-field in earlywood, one in latewood.

- Pujana *et al.* (2014, 2015) described a Palaeocene and Eocene species of *Protophyllocladoxylon* from Antarctica, with tracheid radial pitting uniseriate to biseriate, rarely triseriate, sometimes scattered and mostly contiguous, opposite to alternate, so, of mixed type with pits circular to hexagonal, usually flattened. The axial parenchyma is absent. Cross-field pitting is fenestriform, with 1-2 large oopores (simple pits) per cross-field, so it's very different of our material.

- Vajda *et al.* (2016) described a permineralized conifer wood attributed to *Protophyllocladoxylon* sp., from Korsaröd, southern Sweden, from some volcanoclastic deposits associated to a Jurassic volcanism, and tried palaeo-

climatically the presence of such an exotic taxon in the Sweden Jurassic.

Our specimen has a tracheidoxylic structure, radial pitting of mixed type, usually uniseriate, with round pits, spaced or contiguous (when flattened), parenchyma with thick terminal (horizontal) walls, and uniseriate rays with high biseriations and phyllocladoid cross-fields, with 1, sometimes 2 oopores of phyllocladoid type, inclined, and with large lens-like apertures.

All these details define a special form of *Protophyllocladoxylon*, and analysing them in comparison with other valid species presented above we think we are faced with a new species that we named *Protophyllocladoxylon holbavicum* sp. nov., after the name of the locality area from where it was found. The studied specimen designated holotype has the inventory number 27644 (field number 1040) and is now deposited in G.I.R. Collections, in the National Geological Museum in Bucharest.

Diagnosis of the new species *Protophyllocladoxylon holbavicum* sp. nov.

Secondary wood with tracheidoxylic structure without any canal, with distinct growth-ring boundary, marked by few rows of smaller tracheids. Tangentially seen the tracheids usually unpitted, but sometimes small circular pits in a single vertical row appear. Radial pitting is of mixed type, usually uniseriate with pits circular, spaced or contiguous, with round or oblique elliptic apertures. When contiguous they are slightly flattened. Axial parenchyma few, disperse, with thick terminal (horizontal) walls. Medullary rays uniseriate, the taller ones with high biseriations giving aspect of biseriate rays with long uniseriate endings, radially homocellular, tangential walls are thick-walled or slightly nodular, no indentures and cross-fields with 1-2 phyllocladoid oopores, bordered, inclined, with large lens-like apertures.

Order **Ginkgoales** Gorozhankin, 1904

Family **Ginkgoaceae** Engl. in Engler et Prantl 1897

Genus ***Palaeoginkgoxylon*** Feng, Wang et Rößler, 2010

Palaeoginkgoxylon sp.

Fig. 10 a-i; Fig. 11 a-i.

The studied material

The studied material is represented by three decimetric-sized samples of silicified wood, two found in the Holbav locality area (South Braşov County), on the Maiului brook and one around Cristian locality, on the Fabricii valley, in the Early Jurassic continental deposits maybe with pyroclastic character considered of Sinemurian-Pliensbachian age, most probably Pliensbachian. The studied sample represent trunk or thick-branch fragments, having dark to black color and by magnifying glass or even by naked eye, regular fibrous structure without vessels suggesting a conifer wood. The specimens with field-numbers 1011, 1023, 1017Cr - in "Grădinaru Collection", are deposited now under the inventory numbers 27645 - 27747 in the collections of the Geological Institute of Romania (G.I.R. Collections), hosted by the National Geological Museum in Bucharest.

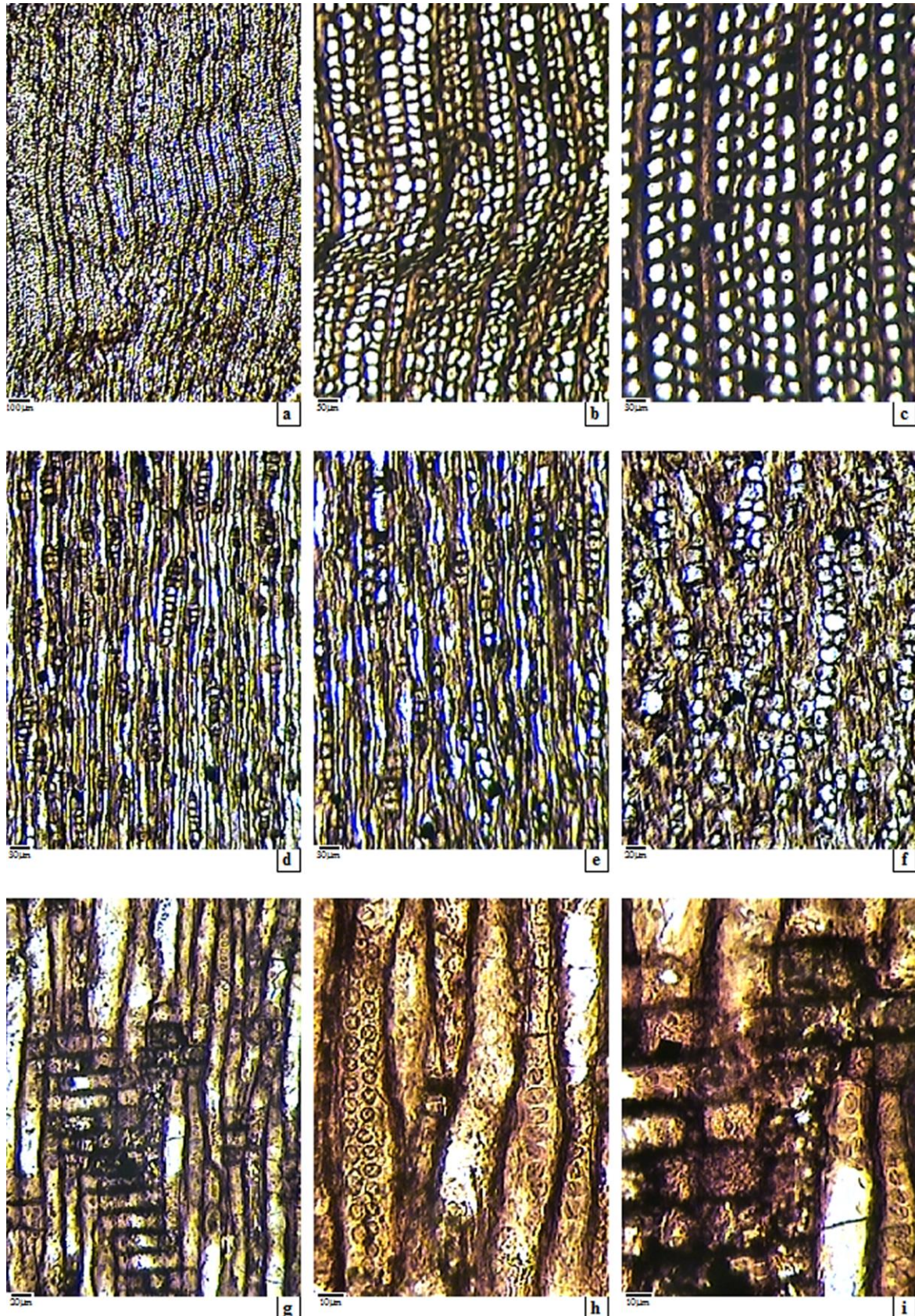


Fig. 10 *Palaeoginkgoxylon* sp. - material - inventory no. 27746 (field no. 1023). **a-c** Cross section: indistinct growth-ring boundary, thick-walled polygonal tracheids with rounded corners, and linear rays. **d-f** Tangential section: unpitted tracheids, uniseriate rays, and biseriate with long endings; polygonal-rounded ray-cells. **g-i** Radial section: radial pitting on tracheids of mixed type, cross-fields with 1-2 rows of oculipores.

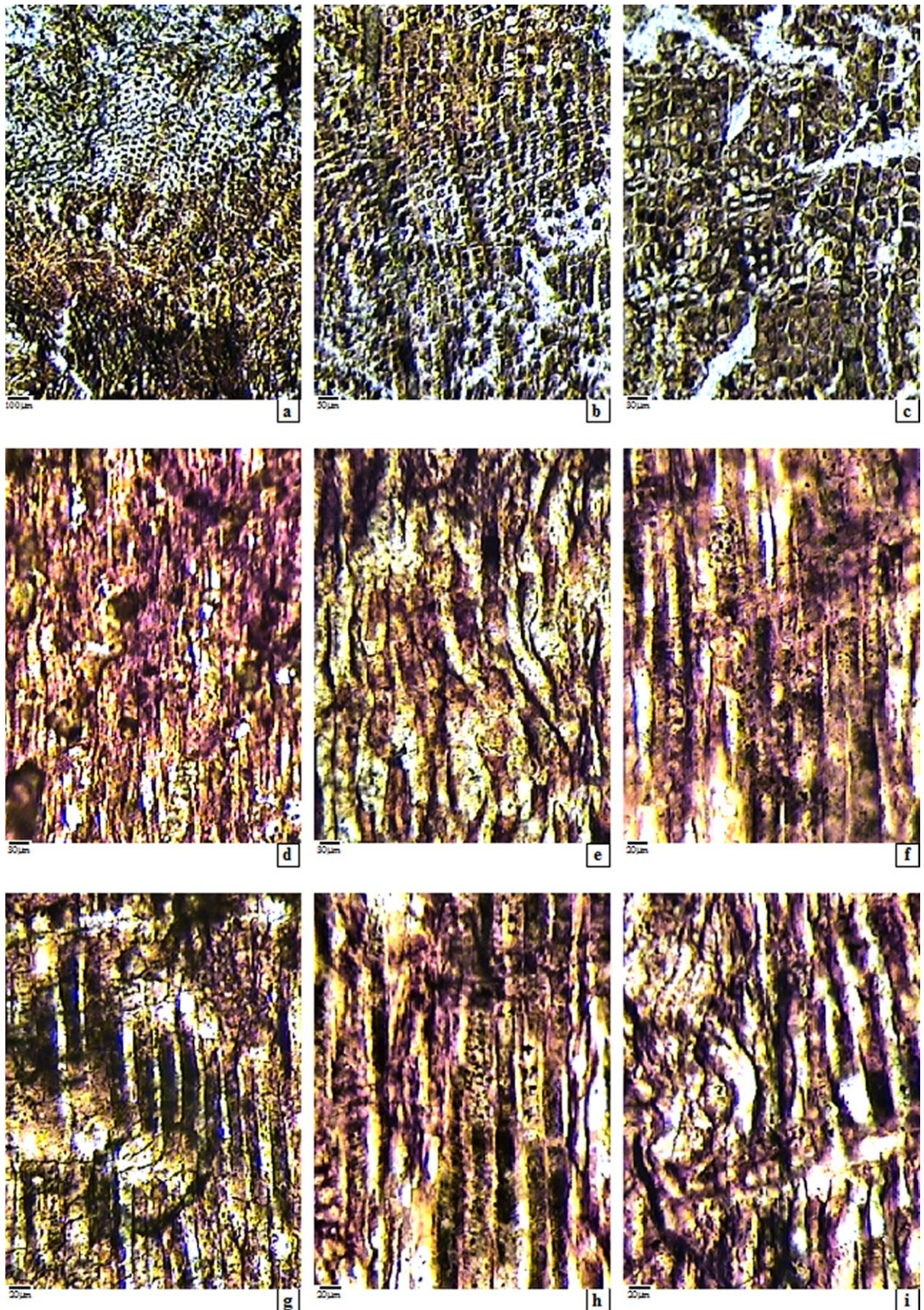


Fig. 11 *Palaeoginkgoxylon* sp. - material - inventory no. 27747 (field no. 1017cr). **a-c** Cross section: less distinct growing boundary, thick-walled polygonal tracheids with rounded corners, and linear rays. **d-e** Tangential section: tracheids with bent ends, uniseriate rays. **f-i** Radial section: tracheids have bent ends which overlap each other, radial pitting on tracheids of mixed type, bad preserved cross-fields.

Microscopic description

Both the studied specimens present secondary wood with tracheidoxylitic structure, with relatively thick-walled tracheids and few parenchyma, devoid of any resin canals, and with quite indistinct growth-rings boundaries, relatively badly preserved (1017).

T r a c h e i d s - with polygonal cross-section with slightly rounded corners, unequal in size, with polygonal or quadrangular rounded lumina gradually diminishing in size from early to late wood, having radial/tangential diameters of 25-35/25(35-40) μm , smaller in the late wood, and not too thin-walled, of 4-7 μm double wall thicker in the late-wood. The interradian bundles have 2-12 radial regular rows of tracheids and the density is 1360-1728 tracheids per mm^2 . Tangentially seen the tracheids are usually unpitted, rarely uniseriate pitted and with the terminal parts slightly leaned over the neighbor cells. Radial pitting is of mixed type, with pits either polygonal to slightly rounded, spaced or contiguous, in one vertical row, or less, biseriate, alternate or slightly irregular to opposite. The bordered pits are polygonal to rounded, slightly flattened, relatively small, of 8-10 μm in diameter when biseriate, and round of 10-16 μm in diameter, also slightly flattened when uniseriate. Often only the pit chamber is visible, the aperture difficult to observe is small round to short elliptic, of 1.5-2(3.5) μm . Crassulae or helical thickenings were not observed. Sometimes the tracheids have bent ends which overlap each other (specimen 1017Cr) and bear some dark content inside lumina.

A x i a l p a r e n c h y m a - appears sometimes like few dispersed cells in transversal view and isolated or in short vertical rows usually, difficult to observe due to bad preservation.

M e d u l l a r y r a y s - are fine, rectilinear in transversal view. Simple pits irregularly arranged could appear on the horizontal walls. Tangentially the usually uniseriate rays have 1-7(-16) cells in height, i.e. 20-180(-350) μm high. Sometimes the taller rays have 1-5 biseriate storeys, giving an almost biseriate aspect (specimen 1023). The ray cells are round to oval and have moderately thin walls, of 2.5-3 μm double walls. Simple pits irregularly present on the tangential walls. Ray density is 7-12 rays on horizontal tangential millimeter. Radially the rays are homocellular, cells all procumbent of 19-20(-28) μm high, moderately thick-walled: 5-6.2 μm the double wall. Within the marginal rows the cells are slightly higher, of 28-30(-40) μm . The cross fields of cupressoid type have 1-4(6-9?) oculipores, rounded or oval of 4-8(13) μm in diameter, with circular or short elliptic tilted apertures of 1-2(3.5) μm . Their arrangement is alternate or slightly irregular on 1-3 rows. Sometimes the walls of ray cells seem to be slightly wrinkled, but few details can be observed, due to the bad preservation.

Affinities and discussions

The presence of some xylotomical details of the secondary wood in our specimens like the presence of bent tracheids' ends overlapping each other and the presence of parenchyma suggested us a possible affinity with some fossil ginkgophyte trunks described from different Mesozoic sites in the world also mainly in the Asian space, in Mesozoic deposits, but in other. As is well known, only a

species of ginkgophyte (*Ginkgo biloba* L.) survived, living now, naturally, in China.

Some fossil taxa show xylotomical details similar to extant *Ginkgo* wood anatomical structure, as *Ginkgoxylon* (Saporta, 1884) Philippe et Bamford, 2008), *Protoginkgoxylon* (Khudajberdyev, 1971) Zheng et Zhang 2008 (in Zheng *et al.*, 2008), *Szeioxylon* Wang, Jiang et Qin, 1994, *Sinopalaeospiroxylon* Zhang et Zheng, 2006 (in Zhang *et al.*, 2006), *Primoginkgoxylon* Süss, Rößler, Boppré et Fischer, 2009; *Palaeoginkgoxylon* Feng, Wang et Rößler, 2010.

Taking into account the diagnosis of the genus *Palaeoginkgoxylon* Feng, Wang et Rößler, 2010 which specify that, beside the primary structure with pith, there is an well developed secondary xylem - which is pycnoxylic (i.e. tracheidoxylitic) with size and slightly irregular arrangement of the tracheid, which, vertically have bent ends overlapping each other, have mixed radial pitting 1-2 seriate, cross-field pits cupressoid, ray cells irregularly pitted on the horizontal and the radial walls and also, and have axial parenchyma - we observed that our studied specimens, even if badly preserved, exhibit more similar features (regarding the tracheidal pitting usually 1-2 seriate, the cupressoid cross fields, the parenchyma), with the secondary wood of the species *Palaeoginkgoxylon zhoui* Feng, Wang et Rößler, 2010, than with other above cited fossil genera.

Thus, based on these similitudes with the type species of a new genus described by Feng *et al.* (2010), we have attribute the studied material (3 specimens) to *Palaeoginkgoxylon* sp., hoping to find a better preserved material and to document, at specific level, the presence of a ginkgophyte wood in the Jurassic flora of the Romanian Carpathians. In this context it's good to remark that Czies (1998a) described *Ginkgo* foliage from the Jurassic of the Carpathian Basin and such identification could support as acceptable Czies's proposal regarding the transfer of *Ginkgoites* and *Baiera* to the genus *Ginkgo*, because in addition to the similar foliage structure of these genera, our Jurassic material xylotomically is very similar the extant *Ginkgo* wood.

CONCLUSIONS

The finding of petrified woods in the Holbav and Cristian areas represents the first and the most significant collection of Early Jurassic lignoflora in the Romanian Carpathians.

The present study is an important contribution to the knowledge on the Mesozoic flora, respectively of the Early Jurassic flora from the Romanian Carpathians. Five Early Jurassic gymnosperm lignotaxa are newly described in the present paper, as follows:

•Order Pinales

- *Agathoxylon holbavicum* sp. nov.
- *Brachyoxylon holbavicum* sp. nov.
- *Brachyoxylon cristianicum* sp. nov.
- *Protophyllocladoxylon holbavicum* sp. nov.

•Order Ginkgoales

- *Palaeoginkgoxylon* sp.

All of the newly described lignotaxa come mainly from Holbav, and these were collected from the lahar deposits in the upper part of the Pliensbachian volcano-

sedimentary sequence cropping out on the Maiului brook. *Brachyoxylon cristianicum* sp. nov. and *Palaeoginkgoxylon* sp. come from the Hettangian-Sinemurian coaly sequence of Cristian, but they also appear in Holbav area. The studied petrified wood material from the Lower Jurassic coaly deposits in the Holbav and Cristian areas may contribute significantly to improve the paleobiogeographic, palaeoenvironmental, paleoecologic and palaeoclimatic reconstructions, already advanced in previous studies of other plant parts (pollen, spores, leaves and fruits imprints).

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E.G. remembers with nostalgia the many years, from 1968 to 1975, when he as a young field-geologist was hosted during summers in the family of Peter and Ida Hull alongside Brigitte, Renate and Peter, in the Vulcan village (Braşov County). E.G. reminds also Stefan Hermanstädter from Vulcan village having a good knowledge on basic mineralogy gained in the Horticulture School before the WWII, who despite his advanced age accompanied him on field and helped to carry by foot several kilometers the fossil woods collected at Holbav on the Maiului brook.

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